

1 **Estimation and long-term trend analysis of surface solar radiation in Antarctica:**

2 **A case study of Zhongshan Station**

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4 Zhaoliang ZENG^{1,2}, Zemin WANG¹, Minghu DING^{2*}, Xiangdong ZHENG², Xiaoyu
5 SUN³, Wei ZHU⁴, Kongju ZHU², Jiachun AN¹, Lin ZANG¹, Jianping GUO², Baojun
6 ZHANG^{1*}

7
8 *¹Chinese Antarctic Center of Surveying and Mapping, Wuhan University, Wuhan*
9 *430079, China*

10 *²State Key Laboratory of Severe Weather, Chinese Academy of Meteorological*
11 *Sciences, Beijing 100081, China*

12 *³National Marine Environmental Forecasting Center, Beijing 100081, China*

13 *⁴Wuhan Economic & Technological Development Zone (Hannan) Meteorological*
14 *Services Wuhan 430079, China*

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29 *Corresponding author: Minghu DING or Baojun ZHANG
30 Email: dingminghu@foxmail.com or bjzhang@whu.edu.cn

ABSTRACT

Long-term, ground-based daily global solar radiation (DGSR) at Zhongshan Station in Antarctica can quantitatively reveal the basic characteristics of Earth's surface radiation balance and validate satellite data for the Antarctic region. The fixed station was established in 1989 and conventional radiation observations started much later in 2008. In this study, a random forest (RF) model for estimating DGSR was developed using ground meteorological observation data, and a high-precision, long-term DGSR dataset was constructed. Then, the trend of DGSR from 1990 to 2019 at Zhongshan Station, Antarctica, was analyzed. The RF model, which performs better than other models, shows a desirable performance of DGSR hindcast estimation with an R^2 of 0.984, root-mean-square error of 1.377 MJ/m², and mean absolute error of 0.828 MJ/m². The trend of DGSR annual anomalies increased during 1990-2004 and then began to decrease after 2004. Note that the maximum value of annual anomalies occurred during approximately 2004/05 and is mainly related to the days with precipitation (especially those related to good weather during the polar day period) at this station. In addition to clouds and water vapor, bad weather conditions (such as snowfall, which can result in low visibility and then decreased sunshine duration and solar radiation) were the other major factors affecting solar radiation at this station. The high-precision, long-term estimated DGSR dataset enables us to further study and understand the role of Antarctica in global climate change and the interactions between snow, ice and atmosphere.

Keywords: meteorological variables, RF model, estimated historical DGSR,

53 long-term trend analysis

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56 **Article Highlights:**

57 ● A thirty-year DGSR dataset, which was produced by combining in situ
58 meteorological observation records with a random forest model, is presented.

59 ● Among the considered models, the RF model shows the best performance for
60 estimating historical DGSR with an R^2 of 0.984, root-mean-square error of 1.377
61 MJ/m^2 , and mean absolute error of 0.828 MJ/m^2 .

62 ● The long-term DGSR trend generally increased during 1990-2004 and then began
63 to decrease after 2004 at Zhongshan Station.

64 ● In addition to clouds and water vapor, abnormal weather in Antarctica (such as
65 fog, blowing snow and snowstorms) was also a major factor affecting solar
66 radiation.

67

68 **1. Introduction**

69 Daily global solar radiation (DGSR, which includes direct and scattered radiation
70 and refers to the total amount of downward shortwave radiation received by the
71 surface each day) is the ultimate source of energy on Earth (Wild, 2009; Wild et al.,
72 2005). Spatiotemporal variations in DGSR determine the climates and environments
73 on the Earth's surface and drive the water, heat and carbon cycles of the Earth system
74 (Che et al., 2005; Wang and Wild, 2016). Polar regions play a vital role in the Earth's
75 surface radiation balance and climate system because there are many important,
76 complex and interacting feedback mechanisms that closely bind the surfaces of polar
77 regions to the global climate system (Bintanja, 1995; Braun and Hock, 2004; Park et
78 al., 2013; Soares et al., 2019; Stanhill and Cohen, 1997). As a result, solar radiation
79 observations and related research in the Antarctic region have received increasing
80 attention (Choi et al., 2019; Ding et al., 2020; Garbe et al., 2020; Zhang et al., 2019).

81 To date, various types of data have been used to study the radiation balance in
82 Antarctica, including reanalysis data, satellite data, and ground station data (Ding et
83 al., 2020; Scott et al., 2017; Stanhill and Cohen, 1997; Yang et al., 2014; Zhang et al.,
84 2016). Scott et al. (2017) used the Clouds and the Earth's Radiant Energy System
85 (CERES) and CALIPSO-CloudSat-CERES-MODIS datasets to study the seasonal
86 changes and spatial distribution of solar net radiation and cloud radiative forcing in
87 southwestern Antarctica from only 2007 to 2010. Zhang et al. (2016) verified the
88 DGSR of six reanalysis datasets by using surface stations around the Antarctic
89 continent and the given deviation value of each reanalysis product. However,

90 satellite-based and reanalysis-based radiation products often face the problem of
91 deviation between the products and trusted ground observations (Jaross and Warner,
92 2008). Stanhill and Cohen (1997) summarized solar radiation measurements across
93 Antarctica (from 12 stations containing 2 to 36 years of data) since 1957. However,
94 the data from most stations used in the above study were very short-term and
95 nonhomogeneous, and it is difficult to gain a full understanding of the long-term
96 characteristics of DGSR in Antarctica (Lacelle et al., 2016; Stanhill and Cohen,
97 1997). In addition, ground-based DGSR observation sites are rare because the special
98 geographical location and harsh natural environment of Antarctica seriously hinder
99 the study of surface radiation balance (Aun et al., 2020).

100 In fact, beyond Antarctica, DGSR observation sites are similarly sparse and
101 uneven across the world due to various problems, such as expensive instruments (He
102 et al., 2018; He and Wang, 2020; Tang et al., 2013, 2011). However, meteorological
103 variables (especially sunshine duration) interact with DGSR, and the number of
104 conventional meteorological observation stations is greater than that of solar radiation
105 stations (Tang et al., 2010; Zeng et al., 2020). Therefore, to obtain long-term,
106 high-density, ground-based solar radiation products, many studies have estimated
107 DGSR from conventional surface meteorological observations using traditional
108 empirical formulas, physical models and machine learning methods (Chen et al.,
109 2013; Huang et al., 2011; Jiang, 2009; Qin et al., 2011; Tang et al., 2018, 2013; Wang
110 et al., 2016). However, these methods have rarely been used in the Antarctic.

111 The Chinese Antarctic Zhongshan Station (69°22'24.76"S, 76°22'14.28"E) is

112 located on the southeast coast of Pritzker Bay in the Lasman Hills of East Antarctica
113 (Figure 1a) (Ai et al., 2019; Chen et al., 2020). Its meteorological observation field is
114 15 m above sea level and approximately 300 m from the nearest coast (Dou et al.,
115 2019; Yu et al., 2017). The local circulation (including valley breezes, land and sea
116 breezes, and katabatic winds) is complicated due to the multifarious landforms and
117 special geographical location. The climate at the station is characterized by low
118 temperature, large temperature difference between winter and summer, low humidity
119 and strong wind; it has the obvious characteristics of an Antarctic continental climate
120 (Yu et al., 2017). Furthermore, as the station is located in front of the Antarctic inland
121 ice sheet, katabatic winds are very obvious, so the DGSR in this area is affected by
122 many factors (Ding et al., 2019). Note that the station was established in 1989, but the
123 DGSR observations began in 2008. Therefore, we aimed to address the lack of
124 radiation data and improve the understanding of Antarctic radiation and its response
125 to global climate change. This study takes the Chinese Zhongshan Station in
126 Antarctica as an example. Based on conventional ground meteorological observations
127 and existing ground radiation observation data, DGSR is estimated using the optimal
128 machine learning method, and then a long-term (~32 year) radiation dataset is
129 obtained. The long-term trend of DGSR and the effects of clouds, water vapor, and
130 visibility on DGSR were also analyzed in this study.

131 **2. Data and methods**

132 **2.1 Meteorological observation data**

133 Meteorological observations began in March 1989, and so far, 32 years of
134 conventional ground meteorological observation data have been accumulated at
135 Zhongshan Station, Antarctica. To ensure the accuracy and quality of observation
136 data, the observation instruments and methods and the accuracy of the ground
137 observation system have been operated in accordance with the ground meteorological
138 observation standards of the China Meteorological Administration [Ground
139 meteorological Observation Standards of China Meteorological Administration].
140 Before February 2002, instrumental observations were recorded manually four times
141 per day (except sunshine duration, which was recorded hourly) at Zhongshan Station;
142 after February 2002, the observation mode changed to an hourly wired telemetry
143 automatic observation system. For more information on the meteorological variables
144 and current sensor types, please see Table 1. All the sensor sampling intervals were
145 changed to 1-min intervals, and the observation data were recorded 24 hours per day.
146 During the instrument replacement period, parallel observations were performed for at
147 least three months. After filtering the abnormal values, homogenization and quality
148 control were also performed by the China Meteorological Administration. Then, the
149 quality controlled and homogeneous observation data were used for data processing
150 and model construction.

151 Previous work aims to establish a virtual network of DGSR datasets using the

152 random forest (RF) model and high-density ground meteorological observations at
153 ~2400 sites across China (Zeng et al., 2020). Due to the spatiotemporal heterogeneity
154 of DGSR, dummy variables including day of year (DOY), latitude, longitude and
155 altitude at all paired stations were also used as input variables in the prediction model.
156 We used the RF model (as the optimal model) and calculated the relative importance
157 of the variables in this study. This study aimed to construct a high-precision,
158 long-term DGSR dataset for Zhongshan Station, Antarctica. Some dummy variables
159 (latitude, longitude and altitude) and some meteorological variables (e.g., land surface
160 temperature, due to the lack of observations at this station) were not used in this
161 study. As suggested by previous studies (Wang et al., 2016; Zeng et al., 2020),
162 meteorological elements that are highly associated with DGSR were selected as the
163 input variables for the machine learning model. These variables include surface
164 pressure (SP), relative humidity (RH), temperature (Tem), wind speed (WS), and
165 sunshine duration (SSD) (corresponding short and full names are shown in Table 1).
166 The dummy variables (i.e., DOY and month) are also used as input variables in the
167 prediction model (similar to Zeng et al. (2020)). DGSR is affected by cloud cover,
168 water vapor, and aerosols before reaching the surface (Che et al., 2005). However,
169 aerosols over the Antarctic are relatively low, so they will not be discussed in this
170 study. In addition, bad weather events (fog, snowfall, blowing snow and snowstorms)
171 with low visibility are frequent at Zhongshan Station and also affect the DGSR. To
172 further analyze the potential causal factors of the DGSR variations, ground vapor
173 pressure (e , which is calculated by temperature and air pressure and represents water

174 vapor content), cloud cover (CF), low cloud cover (LCF) and visibility (Vis) data are
175 also used in this study (CF, LCF and Vis data were collected by manual visual
176 observation).

177 Solar radiation observations started relatively late (March 2008) at Zhongshan
178 Station, although short-term observations and research projects of solar radiation were
179 performed during the periods of January-February 1990 (Wang and Xiong, 1991) and
180 February 1993 to December 1994 (Bian et al., 1998). The solar radiation dataset for
181 2008 to 2020 was first used in this study. The observation site is located in the
182 meteorological field north of the station (as shown in Figure 1b), where the surface is
183 exposed rock from November to February, and there is snow for a short time in other
184 periods, although there is usually almost no snow. A TBQ-2-B pyranometer (Figure
185 1c) was used to measure global solar radiation with a wavelength range of 0.3 to 3 μm
186 and a resolution of hours at this station. The measured signal range of the TBQ-2-B
187 pyranometer is 0-2000 W/m^2 , the output signal is 0-20 mV, and the annual stability is
188 $\pm 2\%$. To ensure the accuracy of observation data, the TBQ-2-B pyranometer passed
189 the verification and calibration of the China Meteorological Administration before
190 installation.

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192 **2.2 Model development**

193 **2.2.1 Data processing and time matching**

194 The meteorological observation data with a data quality code of 0 (passing all

195 quality control checks) were extracted and time matched. The meteorological data
196 observed at 00, 06, 12, and 18 hours (UTC) each day were then averaged to obtain the
197 daily mean values. The DGSR and SSD were obtained as the sums of 24 hours per
198 day. The final available data include conventional meteorological data from March
199 1989 to March 2020 and radiation observations from March 2008 to March 2020.
200 Figure 3 shows the statistical properties (minimum, maximum, mean, and standard
201 deviation) of the variables used for model training, testing and hindcast estimation at
202 Zhongshan Station during 2010-2020. The DGSR (SSD) ranges from 0 (0) to 36.27
203 MJ/m² (21.70 hours), and the annual average value is 10.04 MJ/m² (4.83 hours). The
204 annual average SP, Tem, RH, and WS are 985.08 hPa, -10.16 °C, 56.90%, and 6.49
205 m/s, respectively.

206 **2.2.2 Model building**

207 We evaluated the performances of machine learning models for estimating DGSR at
208 Zhongshan Station, including RF (Chen et al., 2018), Light Gradient Boosting
209 Machine (LightGBM) (Chen et al., 2019; Ke et al., 2017), decision tree (DT)
210 (Quinlan, 1986), back propagation neural network (BPNN) (Wen et al., 2002),
211 eXtreme Gradient Boosting (XGBoost) (Zelteman, 2015), support vector machine
212 (SVM) (Cortes and Vapnik, 1995), multiple linear regression (MLR) (Zelteman,
213 2015) and Adaptive Boosting (AdaBoost) (Wang, 2012) models. The RF model is a
214 widely used machine learning model that has a highly flexible algorithm and the
215 capacity to analyze complex interactions of data classifications with noise or missing
216 values (Chen et al., 2018). The RF model has also been used as a variable selection
217 tool to select the input variables for a final model (Zeng et al., 2020). The RF model
218 uses a bagging method to produce the training dataset. The out-of-the-bag (out-of-bag,
219 OOB) data were used to evaluate the veracity of the regression predicted by the RF

220 model. OOB estimation was an unbiased estimation when the number of trees was
221 sufficient. In common statistical methods, overfitting occurs frequently when the
222 methods have high degrees of freedom. In contrast to other models (e.g., the BPNN
223 and SVM models), the RF model is an ensemble of random trees and basically has no
224 issue of overfitting. LightGBM is a gradient promotion framework based on decision
225 trees and can model complex nonlinear functions. LightGBM has the advantages of
226 distributed and high performance in sorting, classification, and regression (Chen et al.,
227 2019). The DT is a common and extensively researched solution to classification and
228 prediction (Quinlan, 1986). The BPNN is a multilayer feedforward neural network
229 based on a mathematical technique named Bayesian regularization to convert
230 nonlinear regression into “well-posed” problems (Wen et al., 2002). The BPNN is
231 composed of three layers: an input layer (first layer), hidden layer (middle layer), and
232 output layer (last layer). XGBoost is a boosting algorithm with high performance for
233 various regression and classification issues. The XGBoost method requires less
234 training and time for prediction and can improve computing speed and accuracy (Gui
235 et al., 2020). The SVM was developed by Vapnik-Chervonenkis dimension theory
236 and structural analysis of the minimum risk principle. The SVM exhibits a unique
237 advantage in dealing with small-sample problems, nonlinear cases, and
238 high-dimensional pattern recognition problems by its kernel functions (Cortes and
239 Vapnik, 1995). MLR is the regression analysis involving two or more independent
240 variables. The MLR model can intuitively and quickly analyze correlations between
241 multiple variables and dependent variables (Zeltermann, 2015). Hence, MLR has been
242 widely used in social science, economics, and technology. AdaBoost is an excellent
243 boosting algorithm that combines multiple weak classifiers into a strong classifier.
244 The main purpose of AdaBoost is to train different learning devices on the same
245 training set and then combine these devices to construct a stronger final learning
246 device (Wang, 2012).

247 The model with the best estimated performance will be the final model in this
248 study. As in previous work (Zeng, et al., 2020), to obtain the optimal machine
249 learning model, the results of 10-fold cross-validation (10-fold CV) were used to

250 evaluate the model performances with different parameters (the final parameters of
251 the models are shown in Table 2). In 10-fold CV, the matched pairs were partitioned
252 into ten parts in equal proportion, with the first part as the testing subset and the other
253 nine parts as the training subsets. This step was repeated ten times until every subset
254 was tested, and the estimation results (mainly consisting of the coefficient of
255 determination, R^2 , root-mean-square error, RMSE, and mean absolute error, MAE) of
256 the 10 parts were averaged and used as the accuracy of the final model. The accuracy
257 indicators, including the R^2 , RMSE, MAE, and difference (estimated minus
258 observed), were used to assess the capabilities of the machine learning models and
259 then obtain an optimal model (Gui et al., 2020; Zeng et al., 2020). In this study, we
260 used the data from April 2010 to March 2020 for training, the 10-fold CV method for
261 testing, and the period from April 2008 to March 2010 for evaluating the historic
262 estimates.

263 **2.2.3 Model application**

264 The optimal model obtained from the above models was applied to estimate the
265 DGSR using meteorological measurements recorded at Zhongshan Station from
266 March 1989 to February 2020. The time variations in DGSR estimated by the optimal
267 model were then compared with the observed DGSR. Finally, the long-term historical
268 estimated DGSR at Zhongshan Station was analyzed, and then changes in the trend
269 and the possible factors influencing these changes were further investigated.

270 **2.3 Methods for DGSR trend analysis**

271 Least squares regression has been applied to detect the linear trend in DGSR
272 annual anomalies (Guo et al., 2017). Five-year running means of DGSR anomalies
273 have been used to visually display the DGSR trend (Xue et al., 2019). In addition, the

274 sliding trend analysis method has been used to help examine the time nodes of
275 changes because trends often change with the span of the variable calculation period
276 (Che et al., 2019; Gui et al., 2019). According to the method from Gui et al. (2019),
277 we used Student's t tests to detect the robustness of each trend, and the criterion for
278 statistical significance was set at the 95% confidence level. Since the estimated DGSR
279 dataset and meteorological observation dataset have complete records of the whole
280 year for each year from 1990 to 2019, the study period is set as 1990-2019 for the
281 analyses of the monthly variations in meteorological variables and DGSR and the
282 long-term changes in the DGSR trend.

284 **3. Results and discussion**

285 **3.1 Validation and comparison of models**

286 Figure 4a and 4b shows the scatterplots of the fitted model and 10-fold CV model
287 results of the RF model from April 2010 to March 2020 at Zhongshan Station,
288 Antarctica. Compared with those of other models, we found that the fitted and 10-fold
289 CV results of the RF model have higher R^2 values of 0.997 and 0.988 and lower
290 RMSE (MAE) values of 0.547 (1.189) MJ/m^2 and 0.278 (0.648) MJ/m^2 , respectively.
291 To assess the performance of the hindcast estimated by the RF model, the hindcast
292 estimated results from April 2008 to March 2010 are shown in Figure 4c. We found
293 that the hindcast estimated DGSR presented good consistency with the observed
294 DGSR ($R^2 = 0.984$, $\text{RMSE} = 1.377 \text{ MJ/m}^2$, and $\text{MAE}=0.828 \text{ MJ/m}^2$). To further
295 examine the hindcast performance of the DGSR estimated by the RF model, as an

296 example, we selected historical estimated DGSR in 2009 (obtained from hindcast
297 estimated results during April 2008-March 2010) for comparison with the ground
298 observed DGSR. Figure 5a shows the daily time series (Figure 5b depicts the
299 difference) of observed DGSR and estimated DGSR at Zhongshan Station. These two
300 time series are highly consistent with each other, and the higher daily (monthly) mean
301 difference values mainly occur in the summer, especially in the polar day period, but
302 do not exceed ± 5 (± 0.85) MJ/m². Figure 5b also shows that approximately 96.7%
303 (343 days in 2009) of the difference values (observed DGSR minus estimated DGSR)
304 fell within the range of ± 2 MJ/m². The results indicate that the DGSR estimated by
305 the RF model closely fits the observed DGSR. Considering the CV results and the
306 accuracy of the historic estimates, the RF model is highly recommended for DGSR
307 estimation at Zhongshan Station in this study. The relative importance of the variables
308 in the RF model is illustrated in Figure S1. As shown in Figure S1, SSD plays a
309 dominant role in terms of the relative importance in the RF model and accounts for
310 60.3% of the overall importance. This result is consistent with previous studies
311 showing that SSD is significantly correlated with DGSR (Wang et al., 2016; Zeng et
312 al., 2020). The results also indicate two other dominant variables: DOY and Tem
313 (accounting for 16.8% and 16.1% of the overall importance, respectively). These
314 results suggest that DOY (seasonal effects) and Tem are also critical for DGSR
315 estimation.

316 In addition to evaluating the performance of the RF model, we also evaluated the
317 performances of other commonly used machine learning models for estimating DGSR

318 at Zhongshan Station. From the results of the performance comparison in Table 3, the
319 RF model performs better than the other models, and the LightGBM, XGBoost,
320 BPNN, and gradient boosted regression tree (GBRT) models show similar historic
321 estimation abilities, followed by those of the DT and AdBoost models. The SVM and
322 MLR models have the worst performances. In common statistical methods, overfitting
323 frequently occurs when the methods have high degrees of freedom. In contrast to
324 other models, the RF model is an ensemble of random trees and has no issue of
325 overfitting in this study.

326 Overall, these comprehensive results further confirm that the RF model has
327 reliable performance in estimating historical DGSR. We can expect that it will be
328 feasible to reconstruct the historical DGSR based on meteorological observation data
329 and the RF model. Thus, the estimated historical dataset is used to accurately describe
330 the comprehensive characteristics and changes in the long-term trend of DGSR.
331 Therefore, we mainly use the DGSR estimated by the RF model for 1989 to 2020 in
332 the following trend analysis.

333 **3.2 Monthly and annual variations in DGSR**

334 Before analyzing the changes in the DGSR trend, we calculated the monthly and
335 annual characteristics of DGSR estimated by the RF model and the corresponding
336 meteorological variables over the period of 1990-2019 (as shown in Figure 6). The
337 estimated DGSR shows significant monthly and seasonal changes, in which the
338 DGSR during the half-year of summer (October to March) is significantly higher than

339 that during the half-year of winter (April to September), and the higher DGSR values,
340 up to 30 MJ/m², are mainly during the polar day period, while the DGSR values are 0
341 MJ/m² during the polar night period. The monthly average humidity varies from 56%
342 to 61% from 1990 to 2019 and is higher in summer than in other seasons. The
343 monthly mean temperature varies from -2.90 to 0.24 °C in summer and from -14.74 to
344 -16.12 °C in winter. The monthly average temperature is the highest in January (0.24
345 °C) and the lowest in July (-16.18 °C). Variations in surface pressure at Zhongshan
346 Station are characterized by half-year waves. From January to June, there are periods
347 of high pressure, and other months have periods of low pressure. The highest (lowest)
348 monthly average surface pressure occurs in June (October) and is greater than 988
349 hPa (lower than 981 hPa). The wind speed is the highest in winter, followed by that in
350 autumn, spring and summer. The highest (lowest) monthly average wind speed is 7.94
351 m/s (5.01 m/s) in August (January).

352 Zhongshan Station is located on the edge of the Antarctic continent and is near the
353 Antarctic ice sheet. The winds at this station are mainly affected by a combination of
354 the easterly airflow in the northern part of the Antarctic continent, polar cyclones and
355 katabatic winds. Easterly winds prevail over this station year round. In summer
356 (December, January and February), both the Antarctic continental cold high-pressure
357 system and the circumpolar low-pressure zone are weaker, and the smaller pressure
358 gradient between these two synoptic systems induces a lower wind speed. In winter
359 (April, May and June), the prevailing Antarctic continental high-pressure system
360 strengthens, and the circumpolar low-pressure zone moves southward, which causes a

361 larger pressure gradient and thus results in a higher wind speed. Additionally, the
362 monthly variations in air pressure at Zhongshan Station are closely related to these
363 two synoptic systems. Zhongshan Station is covered by snow and ice and the air has
364 lower relative humidity in winter, while the snow and sea ice around this station melt
365 in summer and thus cause increased saturation of water vapor in the air and higher
366 relative humidity. It is noted that the long sunshine duration and strong solar radiation
367 in summer play key roles in the temperature increase at this station.

368 The monthly variation in SSD is basically consistent with that in DGSR, and the
369 R^2 between SSD and DGSR is 0.88, indicating that SSD is the main input variable for
370 DGSR in the machine learning models. Figures S2 and S3 show that the yearly
371 average value of the estimated DGSR coincides with that of the measured DGSR
372 from 2009 to 2019 (and also in 1994), and the differences (estimated DGSR -
373 observed DGSR) were mainly distributed between -0.1 and 0.1 MJ/m². Furthermore,
374 the annual changes in the estimated DGSR and observed SSD trends are highly
375 consistent, suggesting that historical DGSR estimated by the RF model has high
376 accuracy for further analysis (such as annual anomaly trends and sliding trends).

377 **3.3 DGSR trend**

378 Note the continuous presence and absence of solar radiation during summer and
379 winter, respectively, because the peculiar conditions of polar day (beginning on
380 November 23 and ending on January 21 of the following year at Zhongshan Station)
381 and polar night (beginning on May 27 and ending on July 18 each year at this station)

382 occur in Antarctica. In addition, the DGSR values were 0 MJ/m² during the polar
383 night period. Therefore, the DGSR trend analysis was divided into the following three
384 scenarios: all days of the year (annual, scenario 1), the polar day period (scenario 2),
385 and the period of all days of the year except for polar day and polar night (scenario 3).
386 Figure 7a to Figure 7c shows the DGSR anomalies and their trends for the three
387 scenarios. It can be seen that the trend of the DGSR annual anomalies increased
388 during 1990-2004 and then began to decrease after 2004. However, obvious
389 differences in anomaly values exist among the three scenarios. The variation in DGSR
390 anomalies is the greatest during the polar day period (increasing linear trend of 0.175
391 MJ/m²/year and decreasing linear trend of -0.101 MJ/m²/year, which are significant at
392 the 95% confidence level), followed by that during all days of the year (increasing
393 linear trend of 0.039 MJ/m²/year and decreasing linear trend of -0.025 MJ/m²/year,
394 which are significant at the 95% confidence level), and is lowest during the period of
395 all days of the year except for polar day and polar night (increasing linear trend of
396 0.011 MJ/m²/year and decreasing linear trend of -0.001 MJ/m²/year, which are not
397 statistically significant). In general the trend of DGSR is similar to Europe (except for
398 China, with a decreasing trend between 1990 and 2000) during the period of
399 brightening, slightly ascend to the early 2005s, after which it shows a decrease to the
400 present (Che et al., 2005; Ohmura, 2009).

401 The DGSR is strongest during the polar day period, accounting for approximately
402 43% of the annual global solar radiation. The surface of Earth receives more solar
403 radiation because the sun always stays above the horizon during the polar day period.

404 In 1990-2019, the average DGSR during the polar day period varies between 26
405 MJ/m² and 31 MJ/m². However, the annual average DGSR varies from 9.6 MJ/m² to
406 10.8 MJ/m². Therefore, the anomaly values and the range of variation during the polar
407 day period are higher than those during all days of the year. The anomaly values
408 during the period of all days of the year except for polar day and polar night alternate
409 between positive and negative, indicating that the change in total solar radiation
410 during this period basically has no obvious trend. The maximum value of the annual
411 anomalies occurred during approximately 2004/05 and is mainly related to the days
412 with precipitation (such as snowfall, which can result in low visibility and then
413 decreased sunshine duration and solar radiation) at Zhongshan Station in Antarctica.
414 In contrast to scenario 3 (all days of the year except for polar day and polar night),
415 scenario 2 (the polar day period) and scenario 1 (annual) have similar trends in the
416 anomalies (and the sunshine duration during the polar day period accounts for
417 approximately 45% of the total sunshine duration of each year), indicating that the
418 changes in sunshine duration and DGSR during the polar day period play a leading
419 role in the changes in the trend of the DGSR annual anomalies.

420 Based on Student's t tests, the sliding trends of DGSR for all situations are shown
421 in Figure 7d to Figure 7f to present a more comprehensive analysis of the annual
422 trends. Sliding trends were calculated for the three scenarios, starting in each year
423 from 1990 to 2015 and ending in 2019 with increments of at least 5 years. As shown
424 in scenario 1 (all days of the year), the trend of DGSR increased from 1990 to 2003
425 (although an opposite trend was found during approximately 1993 and 1995), then

426 declined sharply after 2004 (especially in scenario 2: the polar day period) and
427 slightly increased in 2012. The sliding trends for scenario 3 (all days of the year
428 except for polar day and polar night) were smaller and relatively stable compared with
429 those for the other scenarios. When the running mean window was longer than 15
430 years, the DGSR trends first increased and then decreased in scenario 3 (all days of
431 the year except for polar day and polar night), and most of the trends were statistically
432 significant.

433 **3.4 The potential impact factors of DGSR**

434 DGSR is affected by cloud cover, water vapor, and aerosols before reaching the
435 surface (Che et al., 2005). However, aerosols are relatively low over the Antarctic, so
436 they will not be discussed in this study. In contrast, bad weather events (fog, snowfall,
437 blowing snow and snowstorms) with low visibility are frequent at Antarctic
438 Zhongshan Station and will also affect the DGSR. Therefore, the effects of cloud
439 fraction, low cloud fraction, ground vapor pressure (e , which represents atmospheric
440 water vapor content), and visibility (which represents bad weather events) on the
441 DGSR at this station are further detailed and explicitly analyzed. Note that the solar
442 radiation is greatly affected by the solar altitude angle in the polar region (which has
443 the phenomena of polar day and polar night). To avoid the effects of these phenomena
444 and analyze the influence of potential factors on solar radiation, we selected only the
445 matched samples of April and September each year for discussion in this study. Here,
446 all the matched samples were divided into five subsets according to cloud cover

447 (0-20%, 20-40%, 40-60%, 60-80% and 80-100%), low cloud cover (0-20%, 20-40%,
448 40-60%, 60-80% and 80-100%) and ground vapor pressure (0-1 hPa, 1-2 hPa, 2-3
449 hPa, 3-4 hPa and >4 hPa), and then the average DGSR was calculated for each subset,
450 and the results are shown in Figure 8a, 8b and 8c, respectively. Similarly, the matched
451 samples were divided into six subsets according to visibility (0-5 km, 5-10 km, 10-15
452 km, 15-20 km, 20-25 km and >25 km), then the average DGSR was calculated for
453 each subset (see Figure 8d for the results).

454 We found that the DGSR significantly decreased (significantly increased) with
455 increasing cloud fraction and ground vapor pressure (visibility). This is because under
456 cloudy conditions, solar radiation reflects back to the top of the atmosphere, reducing
457 the amount of solar radiation reaching the Earth's surface. However, the DGSR did
458 not change much when the low cloud fraction was greater than 40%. We also found
459 that DGSR was generally low during severe weather with low visibility but
460 significantly higher under high visibility conditions. Overall, the DGSR decreases
461 with increasing cloud cover, low cloud cover and ground vapor pressure but increases
462 with increasing visibility. The change in DGSR with low cloud cover is not
463 significant with the change in cloud cover, ground vapor pressure and visibility.
464 Meanwhile, the times series of DGSR, LCF, CF, e, Vis, and SSD and their long-term
465 trends are examined in Fig. S4, respectively. The DGSR shows a small upward trend.
466 Accordingly, SSD and Vis exhibit upward trends (CF, LCF and e exhibit downward
467 trends), implying that the cloud cover, water vapor and abnormal weather (except as
468 solar altitude angle in the polar region) are the common factors that influence the

469 trend of DGSR at this station. To sum up, clouds and water vapor in the atmosphere
470 are the main factors affecting solar radiation. Bad weather conditions, such as fog,
471 blowing snow and snowstorms, are also a major factor affecting solar radiation at
472 Zhongshan Station, Antarctica.

473 **4. Concluding remarks**

474 Based on ground meteorological observation data, an RF model was developed to
475 estimate DGSR, and a high-precision, long-term DGSR dataset was constructed for
476 1989 to 2020 at Zhongshan Station, Antarctica. Long-term trends and the potential
477 impact factors of DGSR were then analyzed in this study. Compared with those of
478 other models, we found that the fitted and 10-fold CV results of the RF model have
479 higher R^2 values and lower RMSE and MAE, and the hindcast estimated DGSR
480 presents good consistency with the observed DGSR ($R^2 = 0.984$, $RMSE = 1.377$
481 MJ/m^2 , and $MAE=0.828 MJ/m^2$). The RF model is better than other models for
482 reconstructing the historical DGSR based on the meteorological observations in this
483 study. The DGSR trends were very consistent in all situations, and DGSR generally
484 increased during 1989-2004 and then began to decrease after 2004. The sliding trend
485 of DGSR in the all days of the year except for polar night period and the polar day
486 period increased from 1990 to 2003 (although an opposite trend was found during
487 approximately 1993 and 1995 for the all days of the year except for polar night
488 period), then declined sharply after 2004 and slightly increased in 2012, while the
489 sliding trends for the period of all days of the year except for polar day and polar
490 night were smaller and relatively stable. The DGSR decreases with increasing cloud
491 cover, low cloud cover and ground vapor pressure but increases with increasing
492 visibility. The results show that clouds and water vapor are the main factors affecting
493 solar radiation in Antarctica. Meanwhile, bad weather conditions, such as fog,
494 blowing snow and snowstorms, are also a major factor affecting solar radiation at
495 Zhongshan Station, Antarctica. Based on the DGSR estimation method in this study,

496 our plan for future work can be divided into two parts: the first part is to construct a
497 virtual DGSR observation network across the Antarctic region, and the second part is
498 to reconstruct historical site-scale DGSR concentrations through this newly
499 constructed virtual DGSR observation network. It is worth noting that some sites in
500 remote areas of Antarctica lack DGSR datasets. Therefore, these high-precision,
501 long-term DGSR datasets can be used to study the radiation balance and the ultimate
502 source of solar energy in Antarctica.

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503 **Acknowledgments**

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508 Institute of China and the Antarctic expeditioners at the Chinese Zhongshan Station.
509 The long-term (March 1989-February 2020) estimated DGSR dataset can be acquired
510 in the Mendeley data repository (data identification number:
511 <http://dx.doi.org/10.17632/2y2mmhzvcx.1>. Direct URL to data:
512 <https://data.mendeley.com/2y2mmhzvcx>).

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REFERENCES

513

- 514 Ai, S., S. Wang, Y. Li, G. Moholdt, C. Zhou, L. Liu, and Y. Yang, 2019:
515 High-precision ice-flow velocities from ground observations on Dalk Glacier,
516 Antarctica. *Polar Sci.*, **19**, 13-23, <https://doi.org/10.1016/j.polar.2018.09.003>.
- 517 Aun, M., K. Lakkala, R. Sanchez, E. Asmi, F. Nollas, O. Meinander, L. Sogacheva, V.
518 D. Bock, A. Arola, and G. d. Leeuw, 2020: Solar UV radiation measurements in
519 Marambio, Antarctica, during years 2017-2019. *Atmos. Chem. Phys.*, **20**,
520 6037-6054, <https://doi.org/10.5194/acp-20-6037-2020>.
- 521 Bian, L., and Coauthors, 1998: A study of radiative features at the Great Wall and
522 Zhongshan Stations of Antarctica, Quarterly journal of applied meteorology, **9**,
523 160-168. (in Chinese)
- 524 Bintanja, R., 1995: The Local Surface Energy Balance of the Ecology Glacier, King
525 George Island, Antarctica: Measurements and modelling. *Antarct. Sci.*, **7** (03),
526 315-325, <https://doi.org/10.1017/S0954102095000435>.
- 527 Braun, M., and R. Hock, 2004: Spatially distributed surface energy balance and
528 ablation modelling on the ice cap of King George Island (Antarctica), *Glob.*
529 *Planet. Change*, **42**, 45-58, <https://doi.org/10.1016/j.gloplacha.2003.11.010>.
- 530 Che, H., Gui, K., Xia, X., Wang, Y., Holben, B.N., Goloub, P., Cuevas-Agulló, E.,
531 Wang, H., Zheng, Y., Zhao, H., Zhang, X., 2019. Large contribution of
532 meteorological factors to inter-decadal changes in regional aerosol optical depth.
533 *Atmos. Chem. Phys.*, **19**, 10497-10523 <https://doi.org/10.5194/acp-2019-360>.
- 534 Che, H., and Coauthors, 2005: Analysis of 40 years of solar radiation data from China,
535 1961-2000. *Geophys. Res. Lett.*, **32**, 1-5.

536 <https://doi.org/10.1029/2004GL022322>.

537 Chen, C., Q. Zhang, Q. Ma, and B. Yu, 2019: LightGBM-PPI: Predicting
538 protein-protein interactions through LightGBM with multi-information fusion.
539 *Chemom. Intell. Lab. Syst.*, **191**, 54-64,
540 <https://doi.org/10.1016/j.chemolab.2019.06.003>.

541 Chen, G., and Coauthors, 2018: Spatiotemporal patterns of PM10 concentrations over
542 China during 2005–2016: A satellite-based estimation using the random forests
543 approach. *Environ. Pollut.*, **242**, 605–613.
544 <https://doi.org/10.1016/j.envpol.2018.07.012>.

545 Chen, J.-L., G.-S. Li, and S.-J. Wu, 2013: Assessing the potential of support vector
546 machine for estimating daily solar radiation using sunshine duration. *Energy*
547 *Convers. Manag.*, **75**, 311-318, <https://doi.org/10.1016/j.enconman.2013.06.034>.

548 Chen, Y., C. Zhou, S. Ai, Q. Liang, L. Zheng, R. Liu, and H. Lei, 2020: Dynamics of
549 Dalk Glacier in East Antarctica Derived from Multisource Satellite Observations
550 Since 2000. *Remote Sens.*, **12**, 1809, <https://doi.org/10.3390/rs12111809>.

551 Choi, T. i., S. J. Kim, J. H. Kim, H. Kwon, and M. A. Lazzara, 2019: Characteristics
552 of Surface Meteorology at Lindsey Islands, Amundsen Sea, West Antarctica. *J.*
553 *Geophys. Res. Atmos.*, **124**, 6294–6306. <https://doi.org/10.1029/2018JD029556>

554 Cortes, C., and V. Vapnik, 1995: Support-vector networks. *Mach. Learn.*, **20**, 273-297,
555 <https://doi.org/10.1023/A:1022627411411>.

556 Ding, M., D. Yang, M. R. van den Broeke, I. Allison, C. Xiao, D. Qin, and B. Huai,
557 2020: The Surface Energy Balance at Panda 1 Station, Princess Elizabeth Land:

558 A Typical Katabatic Wind Region in East Antarctica. *J. Geophys. Res. Atmos.*,
559 **125**. <https://doi.org/10.1029/2019JD030378>.

560 Dou, Y., G. Zuo, X. Chang, and Y. Chen, 2019: A Study of a Standalone Renewable
561 Energy System of the Chinese Zhongshan Station in Antarctica. *Appl. Sci.*,
562 **9**,1968.

563 Garbe, J., T. Albrecht, A. Levermann, J. F. Donges, and R. Winkelmann, 2020: The
564 hysteresis of the Antarctic Ice Sheet. *Nature*, **585**, 538-544.
565 <https://doi.org/10.1038/s41586-020-2727-5>.

566 Gui, K., and Coauthors, 2019: Satellite-derived PM2.5 concentration trends over
567 Eastern China from 1998 to 2016: Relationships to emissions and meteorological
568 parameters. *Environ. Pollut.*, **247**, 1125–1133.
569 <https://doi.org/10.1016/j.envpol.2019.01.056>.

570 Gui, K., and Coauthors, 2020: Construction of a virtual PM2.5 observation network in
571 China based on high-density surface meteorological observations using the
572 Extreme Gradient Boosting model. *Environ. Int.*, **141**, 105801.
573 <https://doi.org/10.1016/j.envint.2020.105801>.

574 Guo, J., and Coauthors, 2017: Declining frequency of summertime local-scale
575 precipitation over eastern China from 1970 to 2010 and its potential link to
576 aerosols. *Geophys. Res. Lett.*, **44**, 5700–5708.
577 <https://doi.org/10.1002/2017GL073533>.

578 He, Y., and K. Wang, 2020. Variability in Direct and Diffuse Solar Radiation Across
579 China From 1958 to 2017. *Geophys. Res. Lett.*, **47**.

580 <https://doi.org/10.1029/2019GL084570>.

581 He, Y., K. Wang, C. Zhou, and M. Wild, 2018: A Revisit of Global Dimming and
582 Brightening Based on the Sunshine Duration. *Geophys. Res. Lett.*, **45**, 4281–
583 4289. <https://doi.org/10.1029/2018GL077424>.

584 Huang, G., M. Ma, S. Liang, S. Liu, and X. Li, 2011: A LUT-based approach to
585 estimate surface solar irradiance by combining MODIS and MTSAT data. *J.*
586 *Geophys. Res. Atmos.*, **116** (D22), <https://doi.org/10.1029/2011JD016120>.

587 Jaross, G., and J. Warner, 2008: Use of Antarctica for validating reflected solar
588 radiation measured by satellite sensors. *J. Geophys. Res. Atmos.*, **113**,
589 <https://doi.org/10.1029/2007JD008835>.

590 Jiang, Y., 2009: Computation of monthly mean daily global solar radiation in China
591 using artificial neural networks and comparison with other empirical models.
592 *Energy*, **34**, 1276-1283, <https://doi.org/10.1016/j.energy.2009.05.009>.

593 Ke, G., and Coauthors, 2017: LightGBM: A highly efficient gradient boosting
594 decision tree, in: *Advances in Neural Information Processing Systems*,
595 3146-3154.

596 Lacelle, D., C. Lapalme, A. F. Davila, W. Pollard, M. Marinova, J. Heldmann, and C.
597 P. McKay, 2016: Solar Radiation and Air and Ground Temperature Relations in
598 the Cold and Hyper-Arid Quartermain Mountains, McMurdo Dry Valleys of
599 Antarctica. *Permafr. Periglac. Process*, **27**, 163-176,
600 <https://doi.org/10.1002/ppp.1859>.

601 Liu, J., F. Weng, Z. Li, and M. C. Cribb, 2019: Hourly PM 2.5 Estimates from a

602 Geostationary Satellite Based on an Ensemble Learning Algorithm and Their
603 Spatiotemporal Patterns over Central East China. *Remote Sens.*, **11**, 2120.

604 Ohmura, A., 2009: Observed decadal variations in surface solar radiation and their
605 causes. *J. Geophys. Res. Atmos.*, **114**, <https://doi.org/10.1029/2008JD011290>,
606 2009.

607 Park, S.-J., T.-J. Choi, and S.-J. Kim, 2013: Heat flux variations over sea ice observed
608 at the coastal area of the sejong station, Antarctica. *Asia-Pacific J. Atmos. Sci.*,
609 **49**, 443-450, <https://doi.org/10.1007/s13143-013-0040-z>.

610 Qin, J., Z. Chen, K. Yang, S. Liang, and W. Tang, 2011: Estimation of monthly-mean
611 daily global solar radiation based on MODIS and TRMM products. *Appl. Energy*,
612 **88**, 2480-2489, <https://doi.org/10.1016/j.apenergy.2011.01.018>.

613 Scott, R. C., D. Lubin, A. M. Vogelmann, and S. Kato, 2017: West antarctic ice sheet
614 cloud cover and surface radiation budget from NASA A-Train satellites. *J. Clim.*,
615 **30**, 6151-6170, <https://doi.org/10.1175/JCLI-D-16-0644.1>.

616 Soares, J., M. Alves, F. N. D. Ribeiro, and G. Codato, 2019: Surface radiation balance
617 and weather conditions on a non-glaciated coastal area in the Antarctic region.
618 *Polar Sci.*, **20**, 117-128, <https://doi.org/10.1016/j.polar.2019.04.001>.

619 Stanhill, G., and S. Cohen, 1997: Recent changes in solar irradiance in Antarctica. *J.*
620 *Clim.*, **10**, 2078-2086,
621 [https://doi.org/10.1175/1520-0442\(1997\)010<2078:RCISII>2.0.CO;2](https://doi.org/10.1175/1520-0442(1997)010<2078:RCISII>2.0.CO;2).

622 Tang, W., K. Yang, J. He, and J. Qin, 2010: Quality control and estimation of global
623 solar radiation in China. *Sol. Energy.*, **84**, 466-475,

624 <https://doi.org/10.1016/j.solener.2010.01.006>.

625 Tang, W., K. Yang, J. Qin, M. Min, and X. Niu, 2018: First Effort for Constructing a
626 Direct Solar Radiation Data Set in China for Solar Energy Applications. *J.*
627 *Geophys. Res. Atmos.*, **123**, 1724-1734, <https://doi.org/10.1002/2017JD028005>

628 Tang, W., K. Yang, J. Qin, C. C. Cheng, and J. He, 2011: Solar radiation trend across
629 China in recent decades: A revisit with quality-controlled data. *Atmos. Chem.*
630 *Phys.*, **11**, 393-406, <https://doi.org/10.5194/acp-11-393-2011>.

631 Tang, W., K. Yang, J. Qin, and M. Min, 2013: Development of a 50-year daily
632 surface solar radiation dataset over China. *Sci. China Earth Sci.*, **56**, 1555-1565,
633 <https://doi.org/10.1007/s11430-012-4542-9>.

634 Wang, G., Xiong X, 1991: Analysis of some characteristics of solar radiation at
635 Zhongshan Station, Antarctica. *Antarctic Research*, **3**, 64-68. (in Chinese)

636 Wang, L., O. Kisi, M. Zounemat-Kermani, G. A. Salazar, Z. Zhu, and W. Gong, 2016:
637 Solar radiation prediction using different techniques: Model evaluation and
638 comparison. *Renew. Sustain. Energy Rev.*, **61**, 384-397,
639 <https://doi.org/10.1016/j.rser.2016.04.024>.

640 Wang, R., 2012: AdaBoost for Feature Selection, Classification and Its Relation with
641 SVM, A Review. *Phys. Procedia*, **25**, 800-807,
642 <https://doi.org/10.1016/j.phpro.2012.03.160>.

643 Wang, Y., and M. Wild, 2016: A new look at solar dimming and brightening in China.
644 *Geophys. Res. Lett.*, **43**, 11,777-11,785. <https://doi.org/10.1002/2016GL071009>

645 Wen J., J. L. Zhao S. W. Luo, and Z. Han, 2002: The improvements of BP neural

646 network learning algorithm. Preprints, *WCC 2000-ICSP 2000. 2000 5th*
647 *international conference on signal processing proceedings. 16th world computer*
648 *congress 2000*, 1647-1649.

649 Wild, M., 2009. Global dimming and brightening: A review. *J. Geophys. Res. Atmos.*,
650 **114**, <https://doi.org/10.1029/2008JD011470>.

651 Wild, M., and Coauthors, 2005: From dimming to brightening: Decadal changes in
652 solar radiation at Earth's surface. *Science*, **308**, 847-850,
653 <https://doi.org/10.1126/science.1103215>.

654 Xue, W., and Coauthors, 2019: Declining diurnal temperature range in the North
655 China Plain related to environmental changes. *Clim. Dyn.*, **52**, 6109–6119.
656 <https://doi.org/10.1007/s00382-018-4505-8>.

657 Yang, Y., S. P. Palm, A. Marshak, D. L. Wu, H. Yu, and Q. Fu, 2014: First
658 satellite-detected perturbations of outgoing longwave radiation associated with
659 blowing snow events over Antarctica. *Geophys. Res. Lett.*, **41**, 730-735,
660 <https://doi.org/10.1002/2013GL058932>.

661 Yu, L., and Coauthors, 2019: The variability of surface radiation fluxes over landfast
662 sea ice near Zhongshan Station, east Antarctica during austral spring. *Int. J. Digit.*
663 *Earth*, **12**, 860-877, <https://doi.org/10.1080/17538947.2017.1304458>.

664 Zelterman, D., 2015: Applied Multivariate Statistics with R. Appl. Multivar. Stat.
665 with R ed. *Springer*, <https://doi.org/10.1007/978-3-319-14093-3>.

666 Zeng, Z., and Coauthors, 2020: Daily Global Solar Radiation in China Estimated
667 From High-Density Meteorological Observations: A Random Forest Model

668 Framework. *Earth Sp. Sci*, **7**, <https://doi.org/10.1029/2019EA001058>.

669 Zhang, T., C. Zhou, and L. Zheng, 2019: Analysis of the temporal–spatial changes in
670 surface radiation budget over the Antarctic sea ice region. *Sci. Total Environ*,
671 **666**, 1134-1150, <https://doi.org/10.1016/j.scitotenv.2019.02.264>.

672 Zhang, X., S. Liang, G. Wang, Y. Yao, B. Jiang, and J. Cheng, 2016: Evaluation of
673 the reanalysis surface incident shortwave radiation products from NCEP,
674 ECMWF, GSFC, and JMA using satellite and surface observations. *Remote Sens.*
675 **8**, <https://doi.org/10.3390/rs8030225>.

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678 **Table 1.** The short name, definition and sensor type of variables.

Variable	Unit	Definition	Sensor type
SP	hPa	Atmospheric pressure	Vaisala PTB220
RH	%	Relative humidity	Vaisala HMP35D
Tem	°C	Surface air temperature	Vaisala HMP35D
WS	m/s	Wind speed	XFY3-1
SSD	h	Sunshine duration	Australia DSU12
GSR	MJ/m ²	Global solar radiation	TBQ-2-B
CF	%	Cloud fraction	-
LCF	%	Low cloud fraction	-
Vis	km	Visibility	-
WVP	hPa	Ground vapor pressure	-
DGSR	MJ/m ²	Sum of global solar radiation over a day	-
DOY	-	Day of year	-
Month	-	Month of year	-

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680 Table 2. The final selected values of the main parameters for each model.

Model name	Parameter	Final value
RF	n_estimators	500
	oob_score	True
	n_jobs	2
LightGBM	n_estimators	3500
	num_leaves	600
	learning_rate	0.05
	max_depth	18
XGBoost	max_depth	16
	learning_rate	0.1
	n_estimators	700
GBRT	n_estimators	750
	Learning_rate	0.5
	Max_depth	3
BP	Solver	adam
	Alpha	1e-5
	hidden-layer-sizes	1000,500,100
SVM	tol	0.000001
AdBoost	n_estimators	500
MLR	--	--
DT	--	--

681

--: Indicates the model parameter value was set as the default.

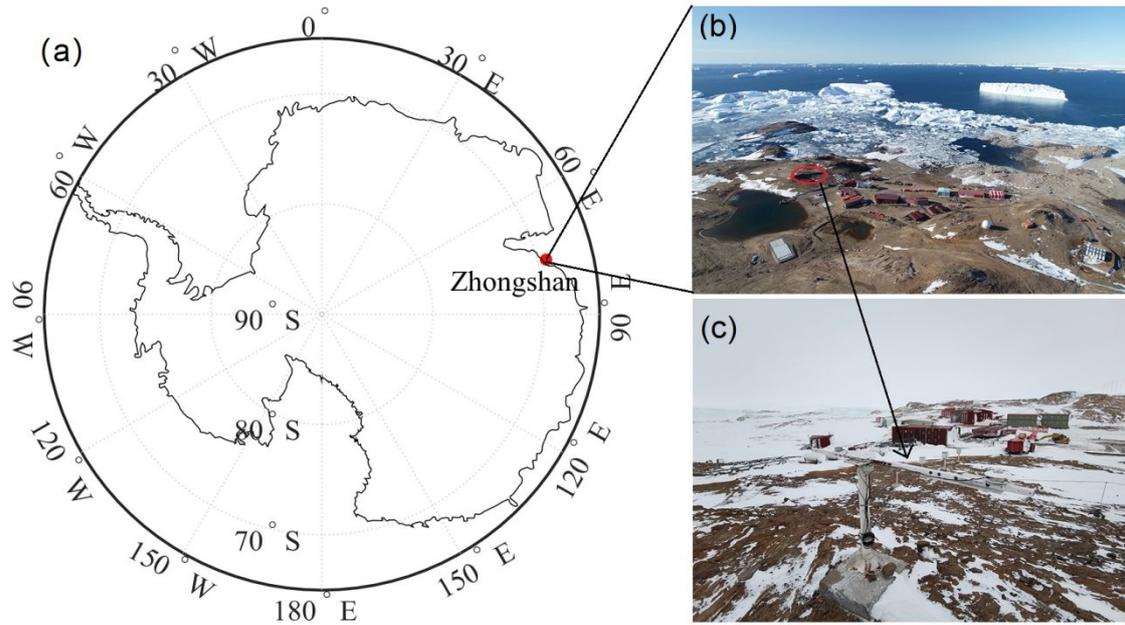
682

683 **Table 3.** The results of the fitted model, 10-fold CV model and historical estimation
 684 power by different machine learning models.

Model name	<u>Fitted model</u>			<u>10-fold CV model</u>			<u>Historical estimation power</u>		
	R ²	RMSE	MAE	R ²	RMSE	MAE	R ²	RMSE	MAE
RF	0.998	0.434	0.238	0.988	1.183	0.648	0.984	1.494	0.846
LightGBM	0.999	0.039	0.027	0.987	1.22	0.78	0.982	1.481	0.884
XGBoost	0.999	0.001	0.013	0.987	1.236	0.678	0.982	1.469	0.845
BP	0.982	1.581	1.132	0.982	1.604	1.151	0.978	1.861	1.370
GBRT	0.999	0.398	0.185	0.98	1.541	0.98	0.974	1.541	0.980
DT	0.998	0.533	0.314	0.976	1.692	0.919	0.967	1.999	1.146
AdBoost	0.962	3.487	3.063	0.959	3.535	3.097	0.961	3.417	3.097
SVM	0.892	5.561	4.755	0.892	5.498	4.687	0.897	3.865	2.980
MLR	0.895	3.498	2.773	0.895	3.500	2.777	0.882	3.778	2.988

685

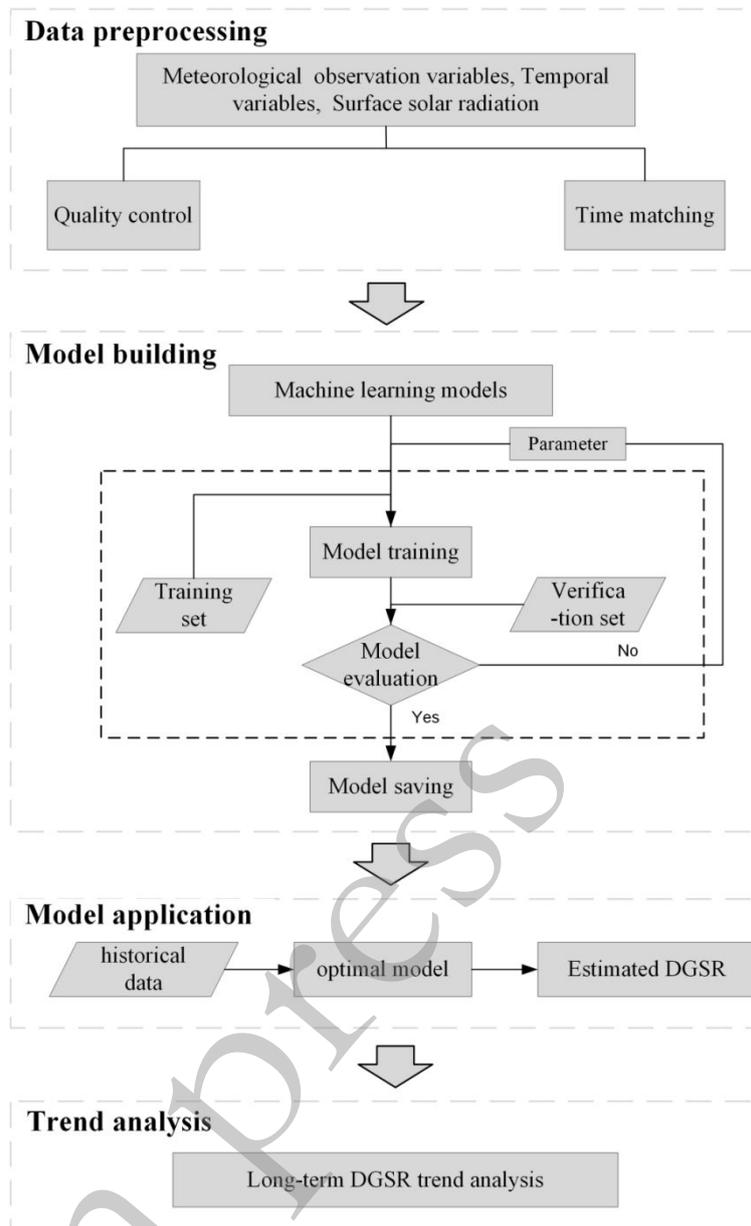
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687 **Fig. 1.** (a) Map showing the location of Zhongshan Station, (b) image of Zhongshan
688 Station area, and (c) the solar radiation instrument in the meteorological observation
689 field.

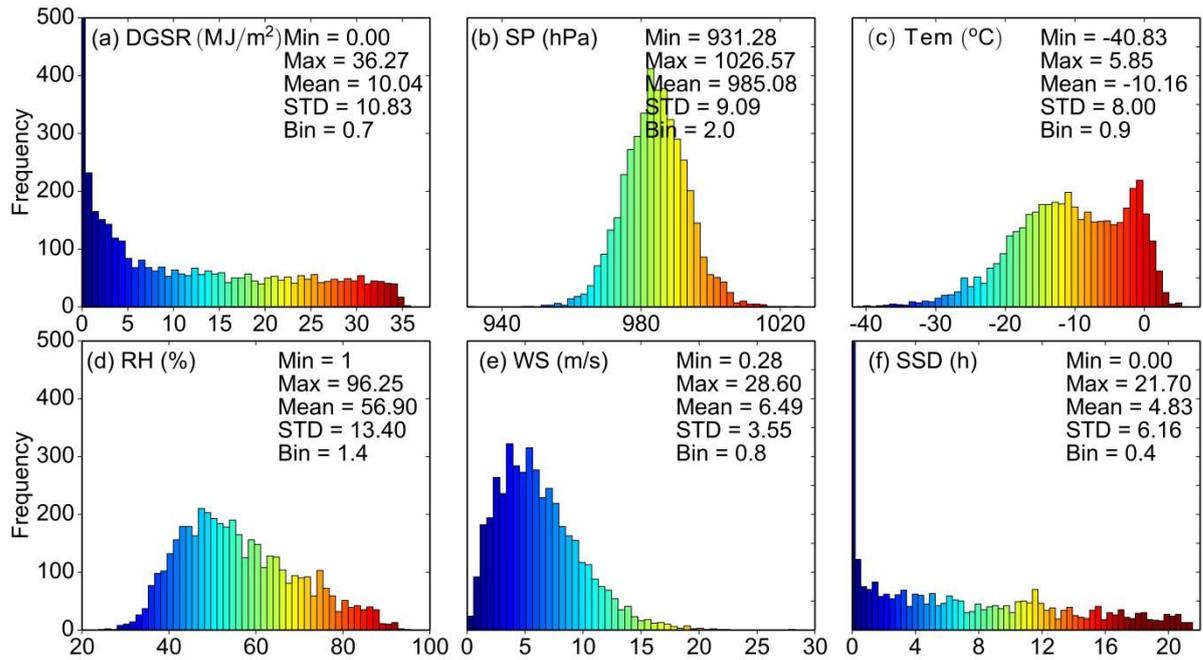
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690

691 **Fig. 2.** Steps of historical estimation and long-term trend analysis of DGSR at

692 Zhongshan Station, Antarctica.



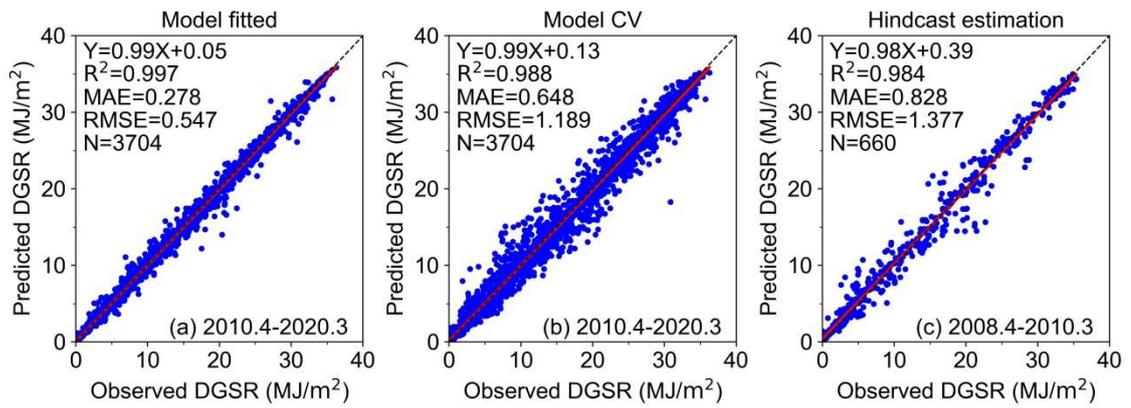
693

694 **Fig. 3.** The frequency distribution of (a) observed DGSR, (b) SP, (c) Tem, (d) RH, (e)

695 WS, and (f) SSD at Zhongshan Station, Antarctica during 2010-2020 for model

696 training and cross-validation.

697



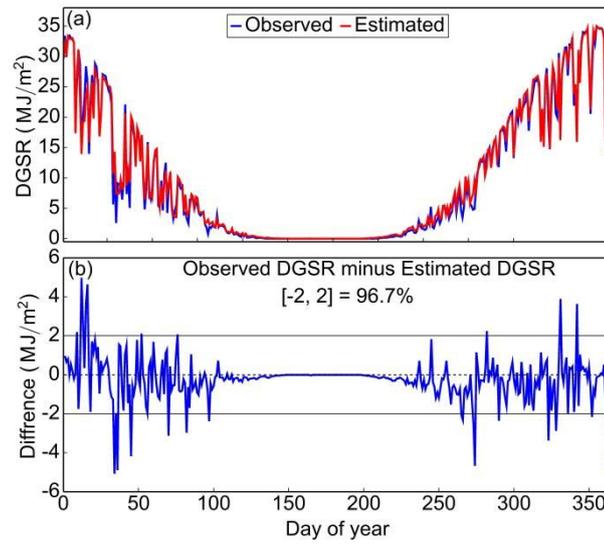
698

699 **Fig. 4.** Scatterplots of the (a) fitted model, (b) CV model and (c) hindcast estimation

700 results of the RF model at Zhongshan Station, Antarctica.

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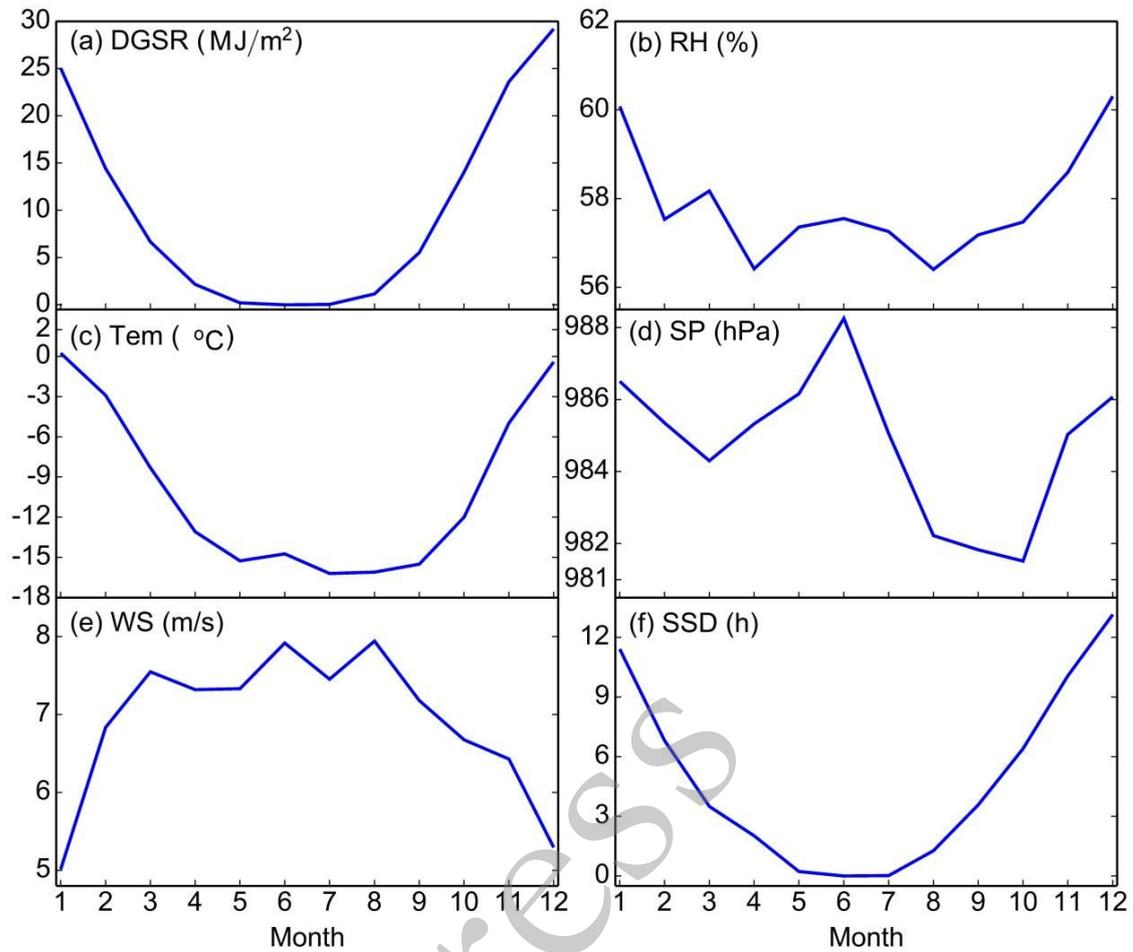
702

703 **Fig. 5.** The (a) time series of observed (blue) versus estimated (red) DGSR and the (b)

704 corresponding difference (observed DGSR - estimated DGSR) in 2009 at Zhongshan

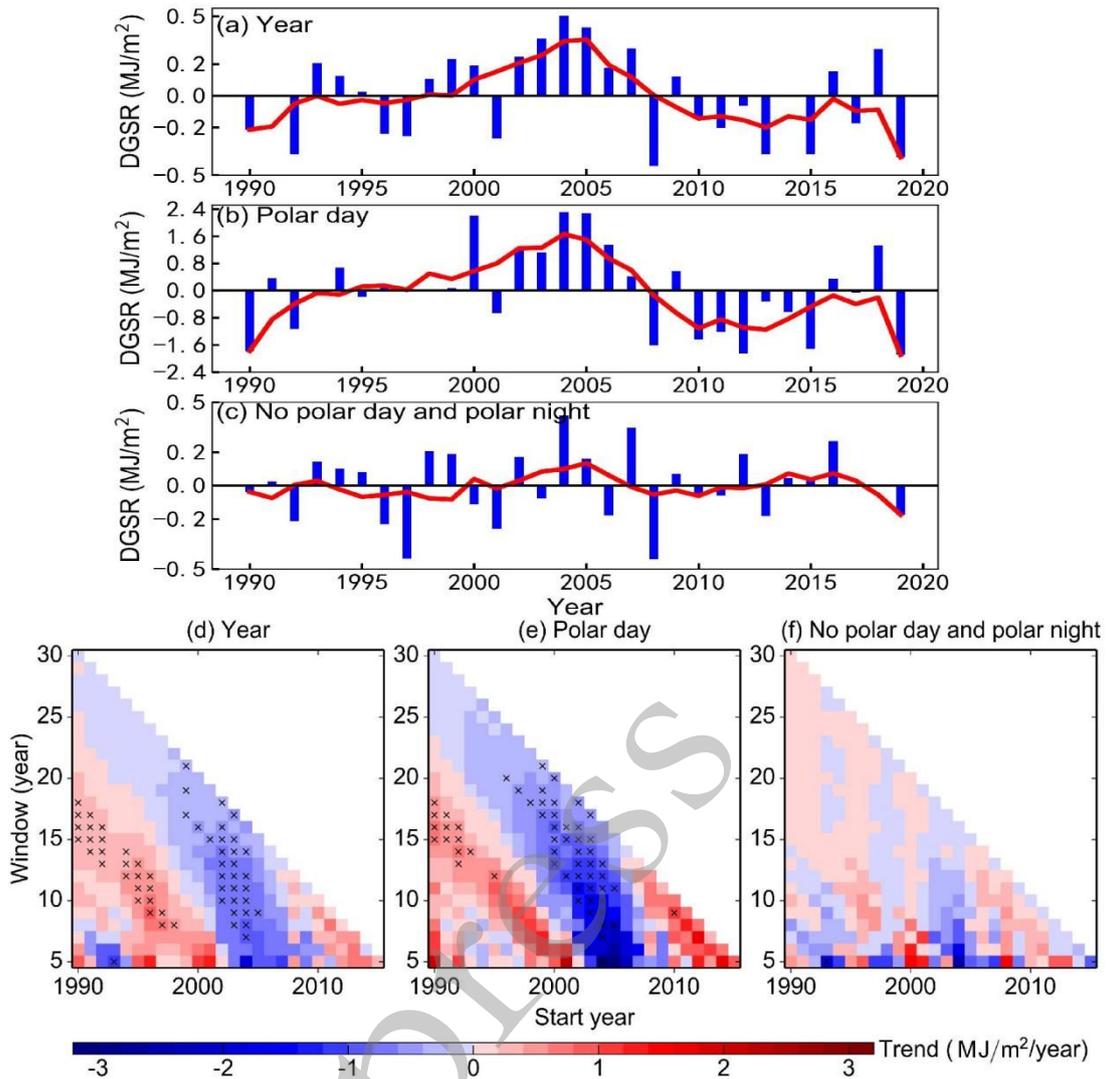
705 Station, Antarctica. [-2, 2] indicates a difference within ± 2 MJ/m².

706



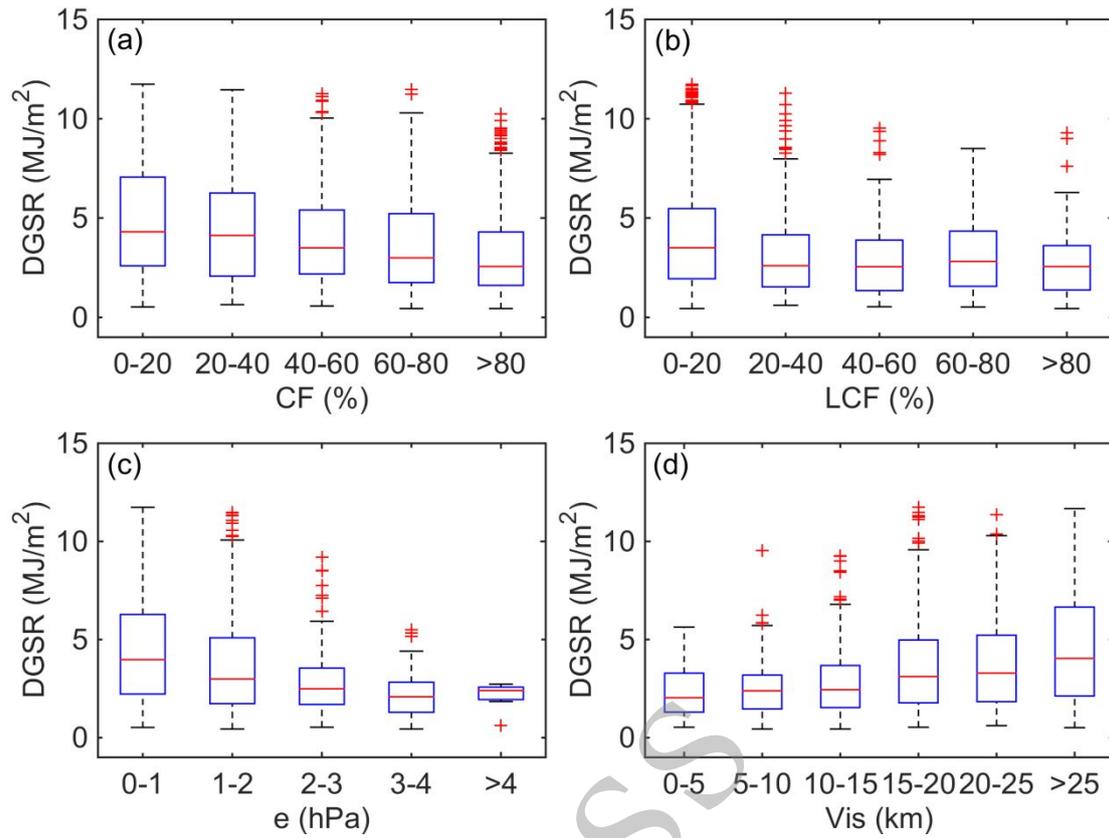
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708 **Fig. 6.** The monthly variation in (a) DGSR (estimated), (b) SP, (c) Tem, (d) RH, (e)
 709 WS, and (f) SSD at Zhongshan Station, Antarctica during 1990-2019.



710

711 **Fig. 7.** Time series of the annual mean anomalies of estimated DGSR: (a) annual
 712 mean but no polar night, (b) polar day and (c) no polar day and no polar night. The
 713 red lines indicate the 5-year running means of the DGSR anomalies. Sliding-window
 714 trend analyses of annual mean estimated DGSR at Zhongshan Station, Antarctica,
 715 from 1990 to 2019 for (d) all years but no polar night, (e) polar day and (f) no polar
 716 day and no polar night.



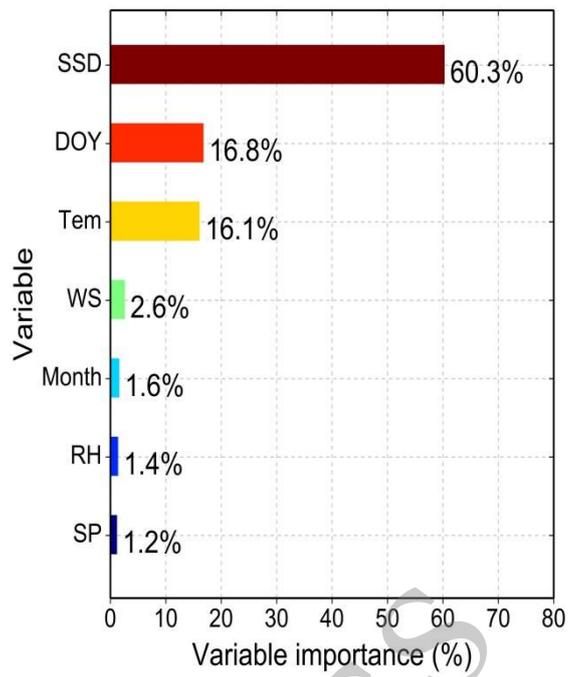
717

718 **Fig. 8.** The effect on the estimated DGSR by the different (a) CF, (b) LCF, (c) e, and

719

(d) Vis conditions at Zhongshan Station, Antarctica.

Supplementary Information

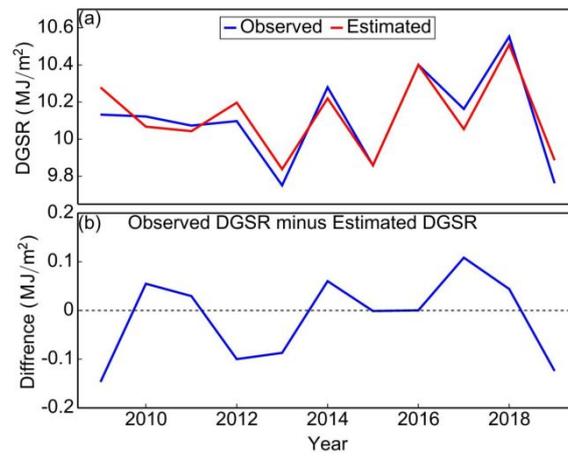


721

722 **Figure S1.** Evaluation of the relative importance of the variables used in the RF
723 models. SSD: sunshine duration; DOY: day of year; Tem: air temperature; WS: wind
724 speed; RH: relative humidity; SP: surface air pressure.

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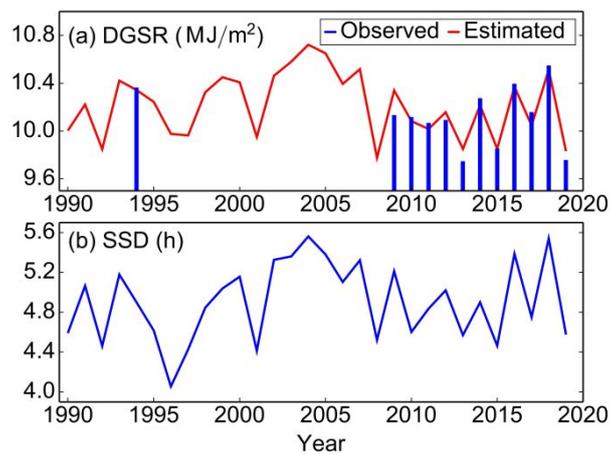
728 Fig. S2. The (a) annual change in observed (blue) versus estimated (red) DGSR and

729 the (b) corresponding difference (observed DGSR - estimated DGSR) from 2009 to

730 2019 at Zhongshan Station, Antarctica.

731

in press



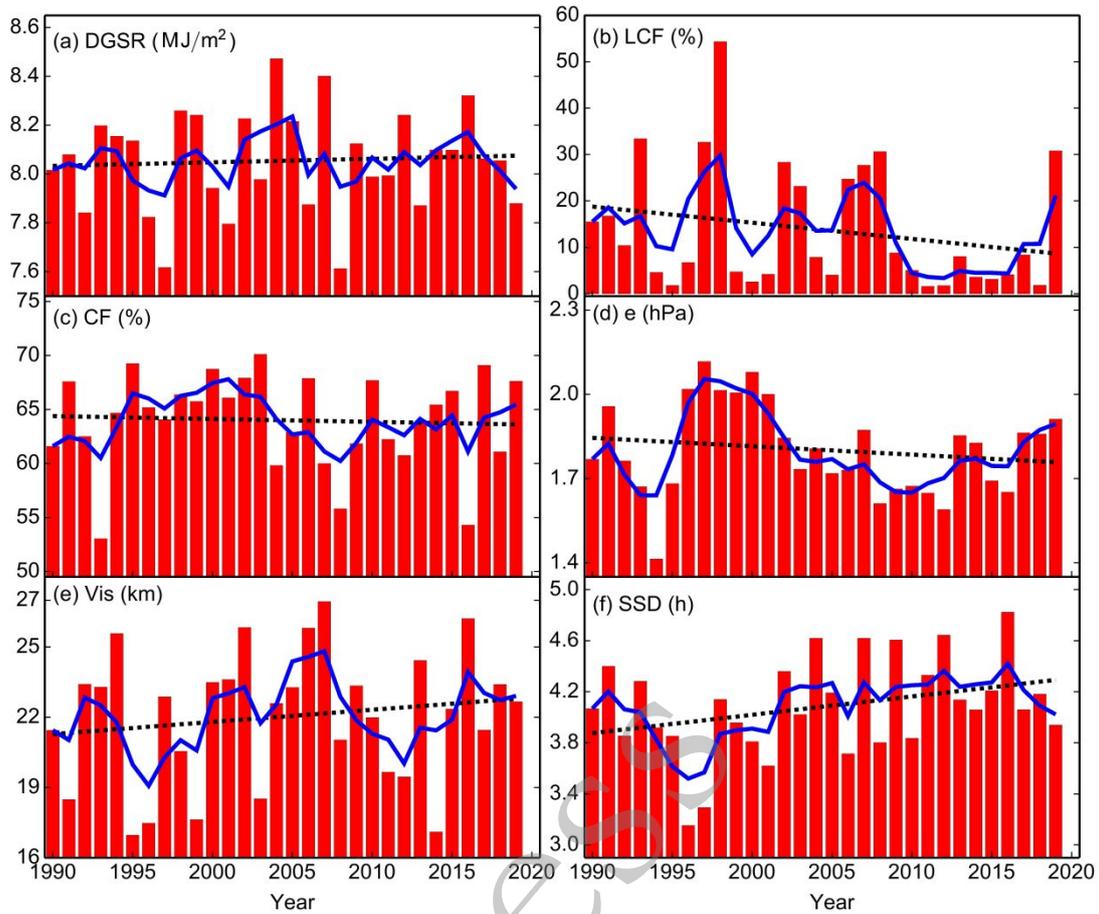
732

733 Fig. S3. The (a) annual change in observed (blue) versus estimated (red) DGSR and

734 the (b) annual change in SSD from 1990 to 2019 at Zhongshan Station, Antarctica.

735

in press



736

737 Fig. S4. The yearly series of (a) DGSR (estimated), (b) LCF, (c) CF, (d) e, (e) Vis,
 738 and (f) SSD at Zhongshan Station, Antarctica from 1990 to 2019 for no polar day and
 739 no polar night. The blue lines indicate their corresponding 5-year running mean
 740 values and straight lines indicate their corresponding linear trends.

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