

1 **The concurrent record-breaking rainfall over Northwest India and North China in**
2 **September 2021**

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10 **ABSTRACT**

11 Extremely heavy rainfall occurred over both Northwest India and North China in
12 September 2021. The precipitation anomalies were 4.1 and 6.2 times interannual standard
13 deviation and broke the record since the observational data were available, i.e., 1901 and
14 1951, respectively. In this month, the Asian upper-tropospheric westerly jet extremely
15 displaced poleward over West Asia, and correspondingly, an anomalous cyclone appeared
16 over India. The anomalous cyclone transported abundant water vapor into Northwest India,
17 leading to the heavy rainfall over there. In addition, the Silk Road pattern, a teleconnection
18 pattern over the Eurasian continent and fueled by the heavy rainfall in Northwest India,
19 contributed to the heavy rainfall in North China. Our study emphasizes the roles of
20 atmospheric teleconnection patterns in concurrent rainfall extremes in the remote regions,

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21 and the occurrence of rainfall extremes during the post- or pre-monsoon in the northern
22 margins of monsoon regions.

23

24 **Key words:** extreme precipitation, Northwest India, North China, westerly jet, Silk Road
25 pattern

26 **Article Highlights:**

27 ● The record-breaking rainfall occurred over both Northwest India and North China in
28 September 2021.

29 ● The rainfall and large-scale circulations in this month resemble the peak rainy season
30 (July and August), possibly due to warmer Eurasian continent.

31 ● The Silk Road pattern is responsible for the concurrence of extremely heavy rainfall
32 over the two remote regions.

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34 **1. Introduction**

35 In September 2021, extremely heavy rainfall occurred in both Northwest India and
36 North China, triggering dire economic and societal consequences in these remote regions.
37 According to India Meteorological Department, in 2021, 8 states and territories in
38 Northwest India recorded excessive precipitation as compared to their respective
39 September averages (<https://weather.com/en-IN/india/monsoon/news/2021-10-01-india-ends-2021-monsoon-season-on-normal-note>). Concurrent with the heavy rainfall in
40

41 Northwest India, precipitation over seven provinces in North China hit a record high since
42 1961 (Liu and Gao, 2021; Zhou et al., 2022b). Wei River, the tributary of Yellow River,
43 with the peak flood in July and August normally, experienced the most severe autumn flood
44 since 1935 (Li et al., 2022). Floods in four provinces affected more than 6 million people,
45 led to 41 people killed or missing, destroyed farmland about 500 thousand hectares, and
46 coasted economic losses totaled 15.34 billion yuan, according to the Chinese government
47 (https://www.mem.gov.cn/xw/yjglbgzdt/202201/t20220123_407199.shtml).

48 The heavy rainfall in September 2021 is unexpected in the climatological sense,
49 because the rainy season is prior to September over both Northwest India and North China.
50 Influenced by the Asian summer monsoon, the rainy season is generally July and August
51 over Northwest India and North China, corresponding to the march of Indian and East
52 Asian summer monsoon, respectively, and subsequently in September, precipitation
53 reduces remarkably (e.g., Krishnamurthy and Shukla, 2000; Wang and Lin, 2002).

54 The seasonal march of rainy season over both regions is closely connected to that of
55 the upper-tropospheric Asian westerly jet (AWJ) (e.g., Liang and Wang, 1998; Ding and
56 Chan, 2005; Chiang et al., 2017; Chowdary et al., 2019; Choudhury et al., 2021; Li et al.,
57 2021). The AWJ is characterized by subseasonal meridional migration due to the seasonal
58 change in warming and resultant meridional temperature gradient (Yeh et al., 1958; Kuang
59 et al., 2007). When the AWJ reaches the northernmost position about 40°N in midsummer,
60 precipitation over Northwest India and North China also reach the peak (e.g.,
61 Krishnamurthy and Shukla, 2000; Ding and Chan, 2005). In addition, the jet-rainfall
62 relationship also exists on the interannual timescale: During the summer when AWJ is
63 displaced poleward (equatorward), the rain bands tend to displaced poleward

64 (equatorward), i.e., more (less) rainfall in Northwest India and North China (Liang and
65 Wang, 1998; Lu, 2004; Wei et al., 2015; Du et al., 2016; Hong et al., 2021; Chowdary et
66 al., 2022).

67 The summer precipitation variations over Northwest India and North China tend to be
68 in phase (e.g., Kripalani and Singh, 1993; Wu, 2017; Wang et al., 2021). This in-phase
69 relationship in summer rainfall can be well explained by the Rossby waves along the AWJ,
70 known as the Silk Road pattern (SRP) (Lu et al., 2002; Wu and Wang, 2002; Saeed et al.,
71 2011; Hong and Lu, 2016; Wu, 2017; Yadav, 2017). The SRP corresponds to anticyclonic
72 (cyclonic) anomalies in the upper troposphere over West Asia and East Asia, respectively,
73 enhancing (suppressing) the rainfall in Northwest India and North China. The Indian
74 rainfall anomalies, affected by the SRP, can in turn trigger the downstream components of
75 SRP, which propagates eastwards along the AWJ (Kripalani et al., 1997; Ding and Wang,
76 2005; Saeed et al., 2011; Greatbatch et al., 2013; Wei et al., 2013; Lin et al., 2017). The
77 SRP can also contribute to the in-phase variations in precipitation between South Asia and
78 East Asia (Wu, 2017; Liu and Huang, 2019). However, it should be mentioned that most
79 of these previous studies focused on summer.

80 In the remnants of this paper, data and methods are presented in Section 2. The extreme
81 rainfall in September 2021 broke the record over both Northwest India and North China is
82 shown in Section 3. We illustrate the circulation anomalies responsible for the extreme
83 rainfall and discuss the linkage of the extreme events between the two remote regions in
84 Section 4. Summary is presented in Section 5.

85 **2. Data and methods**

86 The precipitation over India is from India Meteorological Department, with a
87 $0.25^{\circ} \times 0.25^{\circ}$ resolution for 1901–2021 (Pai et al., 2014). The precipitation over China is
88 from gauge-based observation of 2400 stations for 1951–2021 provided by the National
89 Meteorological Information Center of China
90 (http://101.200.76.197/en/?r=data/detail&dataCode=SURF_CLI_CHN_MUL_DAY_CE
91 [S_V3.0](http://101.200.76.197/en/?r=data/detail&dataCode=SURF_CLI_CHN_MUL_DAY_CE)). The precipitation anomaly over China in Figure 1a was calculated at each station
92 respectively and then interpolated to $0.25^{\circ} \times 0.25^{\circ}$ grid by Cressman interpolation for plot.
93 Wind, specific humidity, 2m air temperature, and vertical integral of water vapor flux are
94 from ERA5 atmospheric reanalysis with a $0.25^{\circ} \times 0.25^{\circ}$ resolution for 1979–2021
95 (Hersbach et al., 2018). The climatological average and interannual standard deviation
96 (STD) are derived by each variable from 1979 to 2020.

97 **3. Extreme rainfall in September 2021**

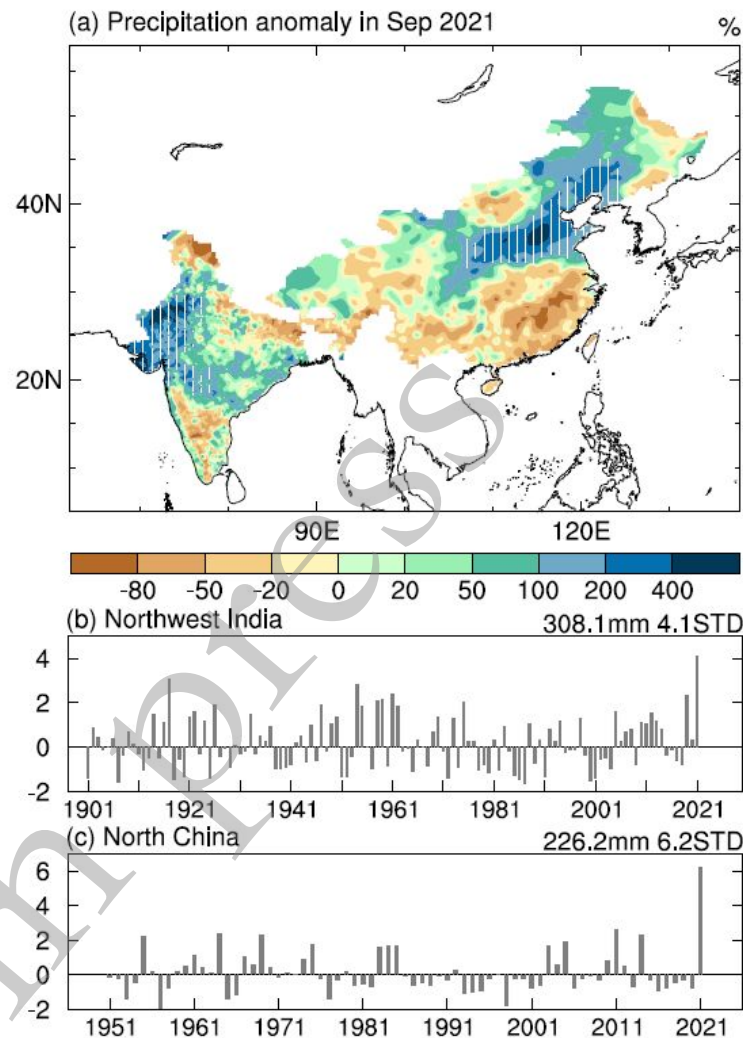
98 Figure 1a shows the precipitation anomalies for September 2021 in percentage. The
99 percentage can highlight the anomalies compared with local climatological-mean
100 precipitation, which exhibits a wide scope, ranging from tiny amount over the deserts over
101 Northwest China to more than 400 mm over the western coastal of Indian subcontinent
102 (Fig. 2b). Generally, rainfall was above normal in the northern parts of both India and
103 China, and below normal in the southern parts. The extremely abnormal precipitation
104 appeared over Northwest India and North China, where the precipitation in September
105 2021 was more than twice the climatological mean, with some areas even more than four
106 times. The rain band appeared over Northwest India and North China in September 2021
107 (Fig. 2a), in a sharp contrast to normal years when the peak rainy season ended in both the
108 regions and accordingly, the rainfall was decreased remarkably (Fig. 2b). Actually, the

109 distribution of rainfall in September 2021 resembles somewhat that for climatological
110 mean in July and August (Fig. 2c), which is the peak rainy season over Northwest India
111 and North China, and thus the precipitation anomalies resemble the subseasonal differences
112 between July/August and September (Fig. 2d).

113 We define Northwest India and North China as the areas where the percentage of
114 precipitation anomalies in September 2021 was higher than 100% within (18° – 30° N, 68° –
115 80° E) and (33° – 45° N, 105° – 125° E), respectively, as shown by the white hatching in Figure
116 1a. The regional-mean precipitation in September 2021 was 308.1 and 226.2 mm for
117 Northwest India and North China, respectively, and it broke the record in both the regions
118 since the observational data are available, i.e., 1901 and 1951, respectively (Figs. 1b and
119 1c). The anomaly over Northwest India was 4.1 times STD, and the second largest dated
120 back to more than 100 years ago (3.1 in 1917). In addition, the anomaly of precipitation
121 over North China was 6.2 times STD, and this value was much greater than the second
122 largest in 2011 (2.6).

123 The correlation coefficient between September precipitation over Northwest India and
124 North China is 0.24 during 1951–2020, significant at the 0.05 confidence level, suggesting
125 that the in-phase relationship in rainfall between Northwest India and North China exists
126 in early autumn, in addition to summer. It should be mentioned that the region definitions
127 for Northwest India and North China are based on the extreme rainfall occurred in
128 September 2021, and thus may not be appropriate for accurately depict the statistical
129 relationship in rainfall between these two regions during a long period. Therefore, the in-
130 phase relationship would be underestimated by the correlation coefficient (0.24), which is
131 calculated by the regional averages. Consider the active role of Indian rainfall, i.e., Indian

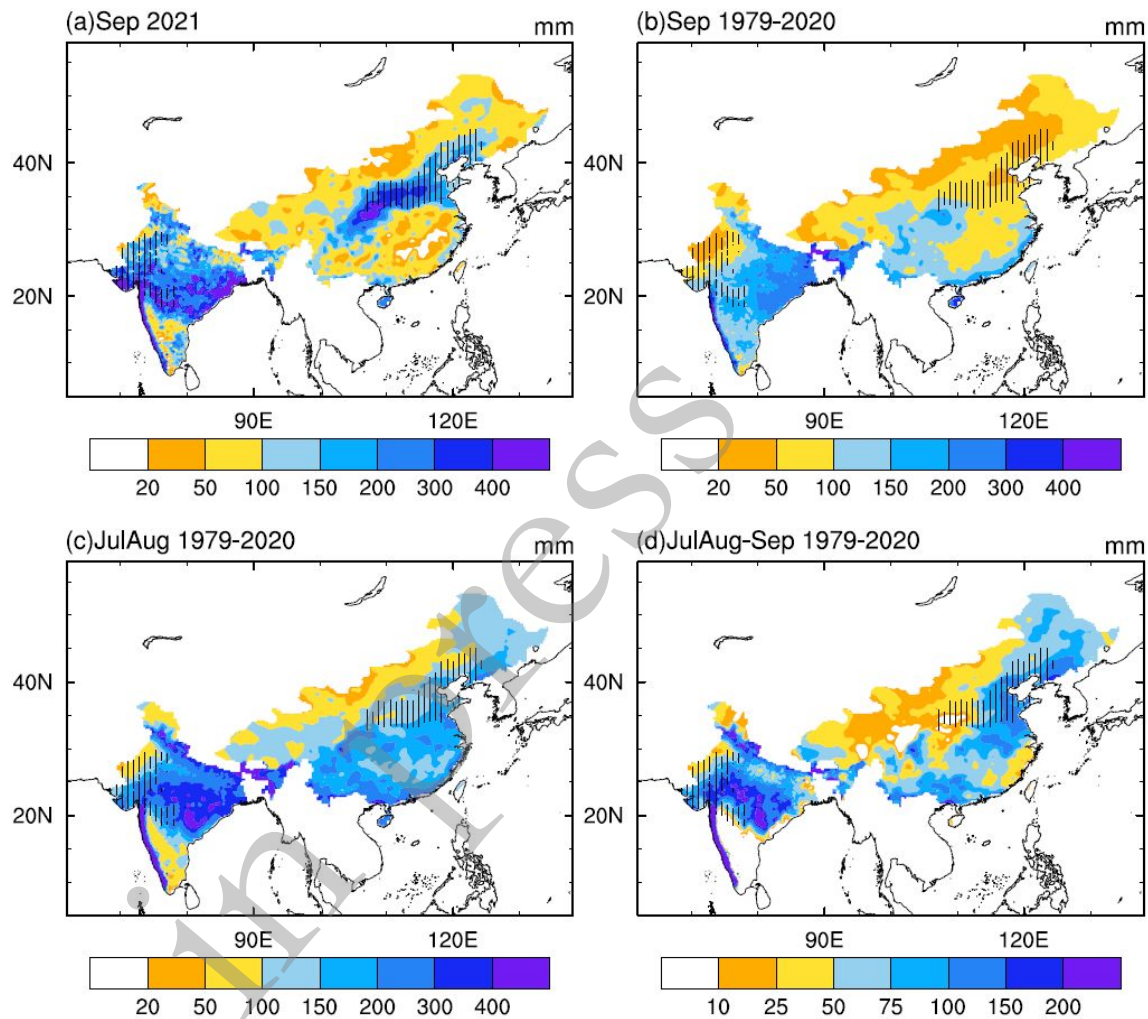
132 rainfall affecting North China rainfall through the SRP, in the following we first discuss
 133 the causes of extreme precipitation over Northwest India in September 2021 and then turn
 134 to North China.



135

136 **Fig. 1.** (a) Percentage of precipitation anomalies (%) in September 2021 relative to
 137 climatology. Only the anomalies at the regions where climatological September
 138 precipitation is higher than 20 mm, except northwest China with sparse stations, are shown.
 139 The white hatching indicates Northwest India and North China. (b) Standardized
 140 precipitation averaged over Northwest India in September from 1901 to 2021. The numbers

141 in top right corner show the average precipitation and times of climatological STD in
 142 September 2021. (c) Same as (b), but for standardized precipitation averaged over North
 143 China.



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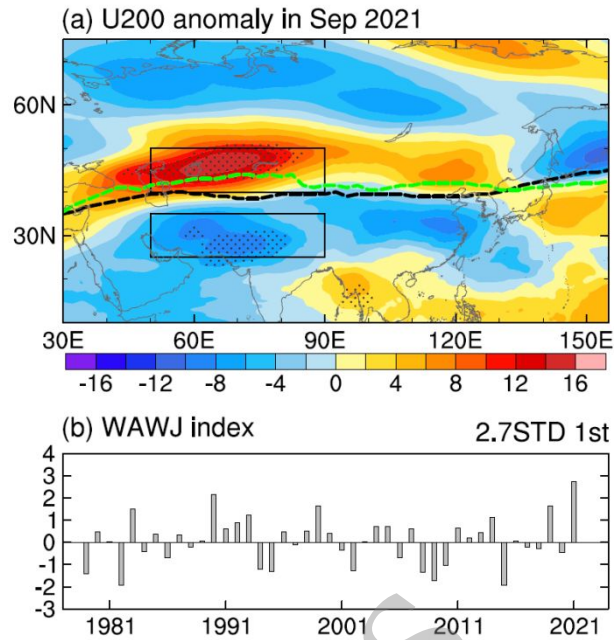
145 **Fig. 2.** Precipitation (mm) (a) in September 2021, (b) averaged in September during 1979–
 146 2020, (c) averaged in July and August during 1979–2020. (d) Difference between the
 147 precipitation averaged in July and August and precipitation in September during 1979–
 148 2020 (mm). Only the contours at the regions where climatological September precipitation
 149 is higher than 20 mm, except northwest China with sparse stations, are shown. The hatching
 150 indicates Northwest India and North China which is same as it in Figure 1.

151

152 **4. Circulation anomalies**153 **4.1 For Northwest Indian rainfall**

154 The 200-hPa zonal wind anomalies in September 2021 are shown in Figure 3a.
155 Westerly and easterly anomalies appeared to the north and south of AWJ axes, respectively.
156 The zonal winds were extremely anomalous over West Asia, and the anomalies exceeded
157 2 times climatological STD at the majority of grids over this region. These zonal wind
158 anomalies corresponded well to the poleward shift of AWJ. In particular, the jet axis (green
159 bold dashed line) shifted poleward about 5 latitude degrees in September 2021 over West
160 Asia in comparison with the climatology (black bold dashed line). Accordingly, the
161 latitudes of the AWJ axis in September 2021 was comparable to, or even higher than, those
162 for the climatology of July and August (Fig. 4b), when the AWJ usually reaches its
163 northernmost location throughout the year, and the zonal wind anomalies in September
164 2021 were stronger than the subseasonal differences between July/August and September
165 (Fig. 4a).

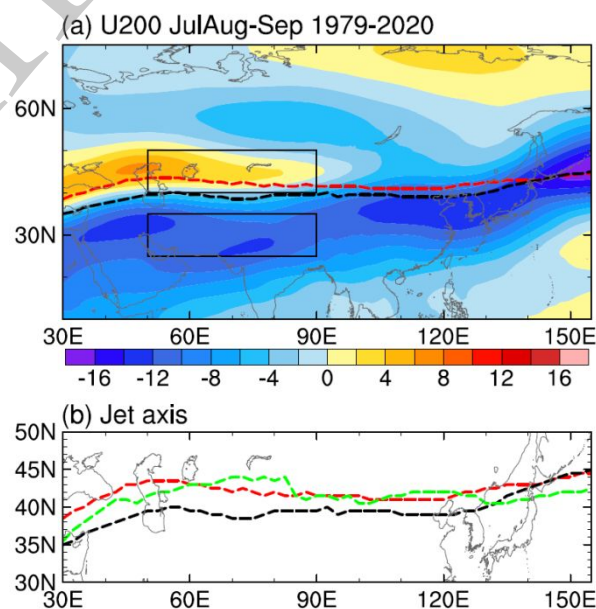
166 Considering the extreme zonal wind anomalies over West Asia, an index is specially
167 defined as the difference between the zonal wind anomalies averaged over (40° – 50° N,
168 50° – 90° E) and (25° – 35° N, 50° – 90° E), to depict the AWJ variability over West Asia
169 (hereafter abbreviated as WAWJ). A positive index represents the poleward displacement
170 of WAWJ, and vice versa. The time series of standardized WAWJ index in September from
171 1979 to 2021 are shown in Figure 3b. The anomaly of WAWJ index in 2021 was 2.7 times
172 climatological STD and was the largest since 1979.



173

174 **Fig. 3.** (a) Anomalies of 200-hPa zonal wind (m s^{-1}) in September 2021. Dots indicate the
 175 anomalies exceeding ± 2 times STD. Green and black dash lines show the AWJ axis in
 176 September 2021 and the climatological AWJ axis, respectively. The black boxes indicate
 177 the regions ($25^{\circ}\text{--}35^{\circ}\text{N}$, $50^{\circ}\text{--}90^{\circ}\text{E}$) and ($40^{\circ}\text{--}50^{\circ}\text{N}$, $50^{\circ}\text{--}90^{\circ}\text{E}$) used to define WAWJ index.

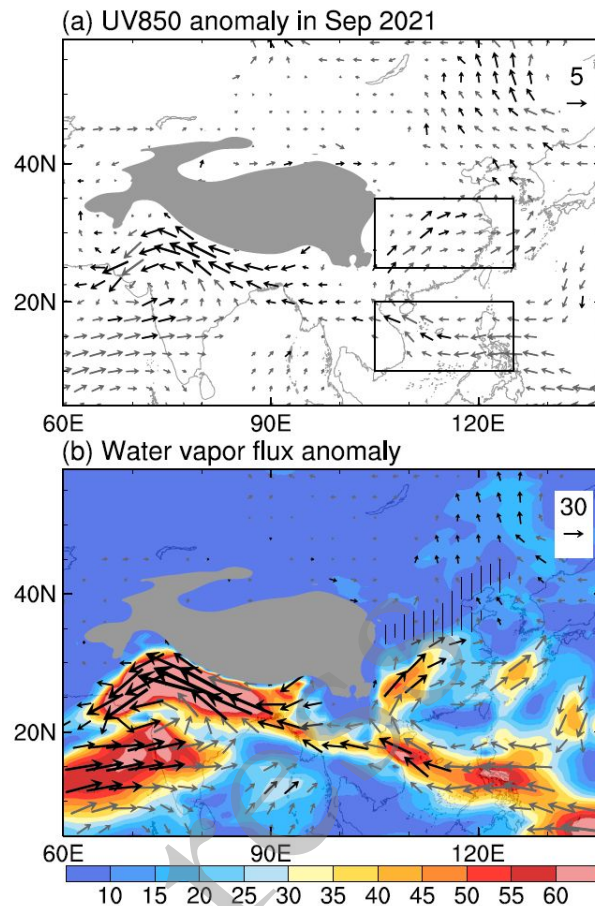
178 (b) Standardized WAWJ index in September from 1979 to 2021. The numbers in right
 179 corner show the anomalous times of STD in September 2021 and the rank of it since 1979.



180

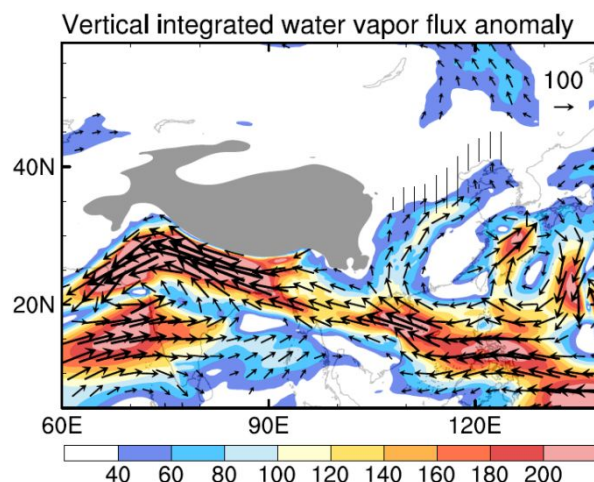
181 **Fig. 4.** (a) Difference of 200-hPa zonal wind (m s^{-1}) between the average of July and August
182 and it in September during 1979–2020. (b) Green, black, and red dash lines show the AWJ
183 axis in September 2021, the climatological AWJ axis in September, and the climatological
184 AWJ axis averaged in July and August, respectively.

185 The lower-tropospheric winds and resultant water vapor flux are crucial for inducing
186 precipitation, and anomalous winds and water vapor flux at 850 hPa (Fig. 5) do explain
187 well the precipitation enhancement in September 2021. There were significant cyclonic
188 anomalies over India and the cyclonic wind anomalies were generally over two times the
189 climatological STD (Fig. 5a). These anomalous winds associated with Indian cyclone
190 transported abundant water vapor at Northwest India (Fig. 5b), consistent with the
191 vertically integrated water vapor flux (Fig. 6), confirming the crucial role of lower-
192 tropospheric winds in transporting total water vapor. It should be mentioned that the
193 monthly circulation anomalies are the sum of various synoptic disturbances including the
194 Cyclone Gulab. This cyclone, which initiated over the Bay of Bengal on September 24 and
195 passed through India afterwards, greatly contributed to the precipitation enhancement over
196 Northwest India (figures not shown).



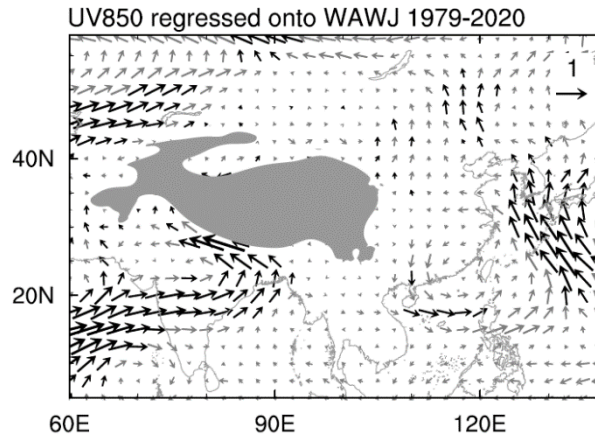
197

198 **Fig. 5.** (a) Anomalies of 850-hPa wind (m s^{-1}) in September 2021. Black and grey vectors
 199 show the wind anomalies exceeding ± 2 and ± 1 times climatological STD, respectively.
 200 The black boxes indicate the regions (25° – 35° N, 105° – 125° E) and (10° – 20° N, 105° –
 201 125° E) used to define anticyclone index in Section 4.2. (b) Anomalies of water vapor flux
 202 ($\text{m s}^{-1} \text{ g kg}^{-1}$) at 850 hPa in September 2021. The contours show the value and the vectors
 203 show the anomalies exceeding ± 2 (black) and ± 1 (grey) times STD. The black hatching
 204 indicates North China which is same as it in Figure 1. The grey shading shows the 2000 m
 205 topography.



206 **Fig. 6.** Anomalies of vertical integrated water vapor flux from the surface to 100 hPa (kg
 207 $\text{m}^{-1} \text{s}^{-1}$) in September 2021. Only the values greater than $40 \text{ kg m}^{-1} \text{s}^{-1}$ are shown. The black
 208 hatching indicates North China which is same as it in Figure 1. The grey shading shows
 209 the 2000 m topography.
 210

211 The concurrent WAWJ poleward displacement in the upper troposphere and Indian
 212 cyclonic anomalies in the lower troposphere may not be accidental. Figure 7 shows the
 213 850-hPa horizontal wind regressed onto the WAWJ index during 1979–2020. There are
 214 significant cyclone anomalies at 850 hPa over India in association with the poleward
 215 displacement of WAWJ, suggesting the close relationship between these circulation
 216 anomalies in the upper and lower troposphere.
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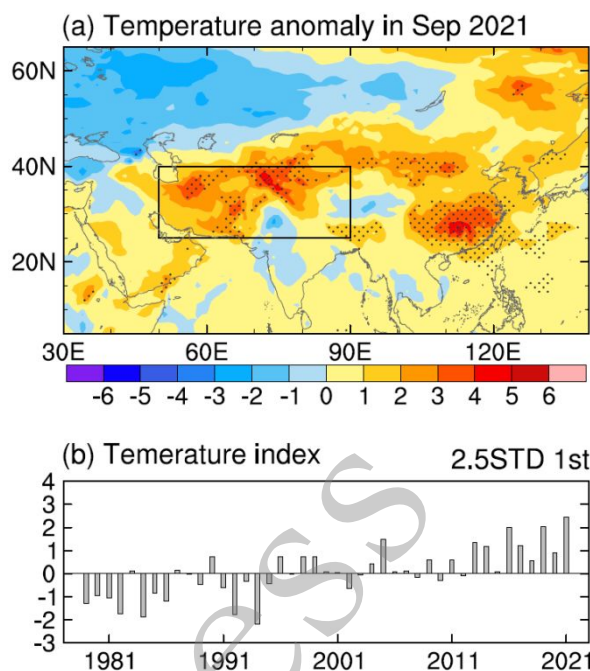


218

219 **Fig. 7.** 850-hPa horizontal wind (m s^{-1}) regressed onto the WAWJ index during 1979–
 220 2020. The black vectors show the wind anomalies significant at the 0.05 level, according
 221 to Student’s test. The grey shading shows the 2000 m topography.

222 A question arises: What is the reason for the poleward displacement of WAWJ in
 223 September 2021? Considering that the meridional temperature gradient plays a crucial role
 224 in determining the meridional migration of AWJ, the surface air temperature anomalies in
 225 September 2021 are shown in Figure 8. The temperature anomalies were extremely positive
 226 over West Asia, exceeding 2 times the climatological STD. On the other hand, the
 227 temperatures were below normal to the north of about 40°N in the Eurasian continent, and
 228 the temperature anomalies were generally weak in the Indian Ocean. As a result, the
 229 meridional temperature gradient was strengthened over the mid-high latitudes and
 230 weakened over the low latitudes in West Asia, corresponding to the poleward displaced
 231 WAWJ. The anomaly of temperature averaged over West Asia (25°–40°N, 50°–90°E) in
 232 September 2021 was the largest since 1979 (Fig. 8b), and the correlation coefficient
 233 between the temperature averaged over this region and the WAWJ index is 0.41 during
 234 1979–2020, significant at the 0.01 confidence level. Thus, the extreme temperature

235 anomalies over West Asia in September 2021 may contribute to the poleward displacement
 236 of WAWJ.



237

238 **Fig. 8.** (a) Surface temperature anomalies in September 2021 (°C). Dots indicate the
 239 anomalies exceeding ± 2 times STD. Black boxes show the region (25° – 40° N, 50° – 90° E)
 240 used for define temperature index. (b) Time series of standardized September temperature
 241 anomalies averaged over the region shown in (a) from 1979 to 2021. The numbers in right
 242 corner show the anomalous times of STD in September 2021 and the rank of it since 1979.

243 **4.2 For North China**

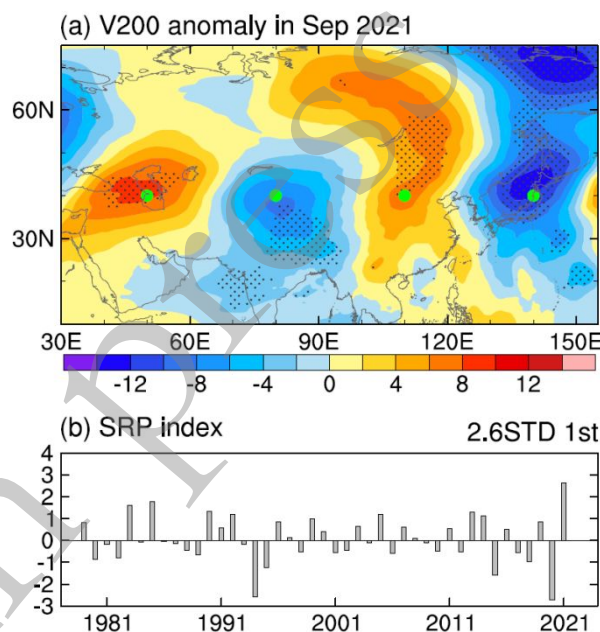
244 Over East Asia, in September 2021 there were also westerly anomalies to the north of
 245 jet axis and easterly anomalies to the south, corresponding to the poleward displacement
 246 of jet axis (Fig. 3a), although these anomalies were not so extreme as those over West Asia.
 247 The latitude of the jet axis in September 2021 was also comparable to, or even higher than,

248 the climatological location of July–August jet axis (Fig. 4b). As mentioned in the
249 introduction, the poleward shifted jet would favor the heavier rainfall in North China.

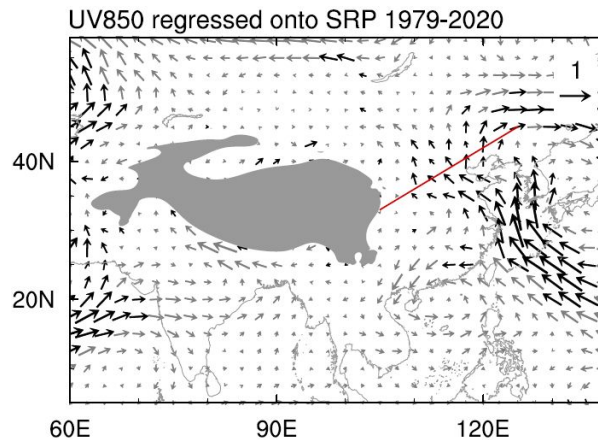
250 The lower-tropospheric circulation anomalies were in favor of water vapor
251 transportation into North China in September 2021 (Fig. 5). Extremely anomalous
252 southwesterlies in association with the anticyclonic anomaly over South China and the
253 western North Pacific transported more water vapor into the western part of North China,
254 and the southerly or southeasterly anomalies over the Yellow Sea and coastal regions of
255 China favored more water vapor transportation into the eastern part of North China. The
256 combination of the anomalous southwesterlies and southeasterlies induced the extreme
257 precipitation over North China. It should be mentioned that the southerly/southeasterly
258 anomalies, albeit much weaker than the southwesterly anomalies, were crucial for water
259 vapor convergence over North China (Figs. 5b and 6), due to the sharp south-to-north
260 decrease of specific humidity over eastern China in September.

261 The southerly and southeasterly anomalies at 850 hPa over the coastal regions of China
262 and Yellow Sea correspond to southerly anomalies in the upper troposphere. In September
263 2021 there were southerly anomalies at 200 hPa over eastern China, and there were also
264 strong southerly and northerly anomalies over the midlatitude Eurasian continent and Japan
265 (Fig. 9a). These meridional wind anomalies resembled well the SRP, and the SRP in
266 September 2021 was extremely anomalous and broke the record (Fig. 9b). Here, the SRP
267 index is defined as the algebraic sum of 200-hPa meridional wind anomalies at the four
268 centers along 40°N as marked in Figure 9a, i.e., $V_{200}(50^{\circ}\text{E}) - V_{200}(80^{\circ}\text{E}) + V_{200}(110^{\circ}\text{E})$
269 $- V_{200}(140^{\circ}\text{E})$.

270 The concurrence between the SRP and lower-tropospheric southerly/southeasterly
 271 anomalies over the coastal regions of China and Yellow Sea is not limited to the specific
 272 case of September 2021, but also exists in normal years. Figure 10 shows the 850-hPa
 273 horizontal wind (m s^{-1}) regressed onto the SRP index during 1979–2020. Stronger SRP
 274 corresponds to significant southeasterly anomalies over the Yellow Sea and northern part
 275 of East China, which leads to the wind convergence over North China. Therefore, the
 276 extremely anomalous SRP in September 2021 was associated with anomalous lower-
 277 tropospheric southeasterlies and wind convergence over North China.



278
 279 **Fig. 9.** (a) Anomalies of 200-hPa meridional wind (m s^{-1}) in September 2021. Dots indicate
 280 the anomalies exceeding ± 2 times STD. Green dots indicate the points (40°N , 50°E),
 281 (40°N , 80°E), (40°N , 110°E), and (40°N , 140°E) used to define SRP index. (b)
 282 Standardized SRP index in September from 1979 to 2021. The numbers in right corner
 283 show the anomalous times of STD in September 2021 and the rank of it since 1979.



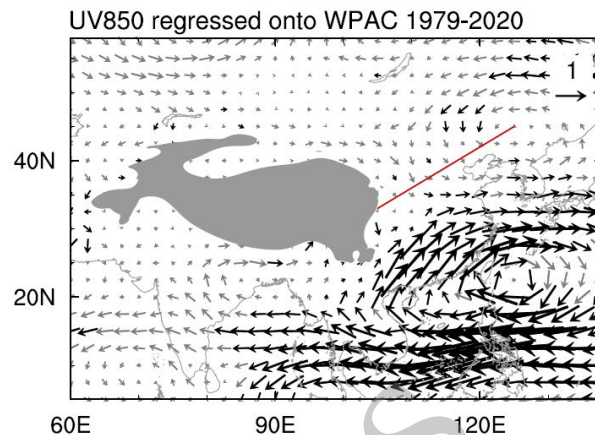
284

285 **Fig. 10.** 850-hPa horizontal wind (m s^{-1}) regressed onto the SRP index (SRPI) during 1979–
 286 2020. The black vectors show the wind anomalies significant at the 0.05 level, according
 287 to Student’s test. The grey shading shows the 2000 m topography. The red line (from 33°N,
 288 105°E to 45°N, 125°E) illustrates the rainfall enhancement over North China in September
 289 2021.

290 The SRP is crucial for the linkage between circulation and precipitation anomalies
 291 over West and East Asia. Hong et al. (2021) stated that the poleward displacement of
 292 WAWJ is associated with the enhanced SRP, and the SRP can be significantly fueled by
 293 the precipitation over Northwest India through latent heat release (Ding and Wang, 2005).
 294 Our results show the SRP can lead to the enhanced precipitation over North China. As a
 295 result, the SRP is responsible for the concurrence of extremely heavy rainfall over the two
 296 remote regions in September 2021.

297 On the other hand, the anomalous anticyclone over South China and the western North
 298 Pacific in September 2021 might be an independent factor. As shown in Fig. 5, the
 299 southwesterly anomalies associated with the anticyclonic anomaly transported more water
 300 vapor northward from ocean in this month. During 1979–2020, it shows weak or even
 301 contrast relationships to the WAWJ displacement (Fig. 7), SRP (Fig. 10), and lower-

302 tropospheric cyclonic anomaly over India (Fig. 11). Therefore, we can conclude that both
 303 this anomalous anticyclone and the SRP contributed to the extremely heavy rainfall in
 304 North China in September 2021.



305

306 **Fig. 11.** 850-hPa horizontal wind (m s^{-1}) regressed onto the index of anomalous
 307 anticyclone over west Pacific (WPAC) during 1979–2020. Here, the index is defined as
 308 the difference between the 850-hPa zonal wind anomalies averaged over (25° – 35° N, 105° –
 309 125° E) and (10° – 20° N, 105° – 125° E), where the winds were extremely anomalous in
 310 September 2021 as shown in Figure 5a. The black vectors show the wind anomalies
 311 significant at the 0.05 level, according to Student’s test. The grey shading shows the 2000
 312 m topography. The red line (from 33° N, 105° E to 45° N, 125° E) illustrates the rainfall
 313 enhancement over North China in September 2021.

314 Recently, Liu et al. (2022) investigated the effects of tropical SST anomalies on the
 315 extreme rainfall event over North China in September 2021. They suggested that the SST
 316 anomalies in the tropical Indian Ocean, Pacific and Atlantic strengthened the convection
 317 over the Maritime Continent and northern Indian Peninsula which contributed to the
 318 anticyclonic anomaly over the western North Pacific and resulted extreme rainfall over
 319 North China. This study, however, emphasizes the role of internal atmospheric variability,

320 i.e., the midlatitude atmospheric circulations, including the SRP and related lower-
321 tropospheric southerlies or southeasterlies.

322 **5. Summary**

323 Record-breaking precipitation occurred simultaneously over Northwest India and
324 North China in September 2021. The precipitation extremes over Northwest India was
325 contributed by the extremely poleward displaced WAWJ and corresponding water vapor
326 transported by the anomalous Indian cyclone. The WAWJ displaced to the northernmost
327 position which might be contributed by the extremely warm temperature over West Asia
328 and resultant meridional temperature gradient. Furthermore, the SRP was extremely
329 anomalous in the month, possibly induced by the Indian rainfall anomalies. The anomalous
330 SRP favored the extreme rainfall in North China through the poleward displaced upper-
331 tropospheric westerly jet over East Asia and the lower-tropospheric southeasterly
332 anomalies, in combination with the water vapor transported by the anomalous anticyclone
333 over the western North Pacific.

334 The present study indicated that the precipitation and circulation in September 2021
335 resemble those in the peak rainy season, i.e., July and August. This may be contributed by
336 the extremely warm temperatures in the Eurasian continent, which could induce the
337 poleward displacement of upper-tropospheric westerly jet (e.g., Lu et al., 2007; Seidel et
338 al., 2008; Pena-Ortiz et al., 2013; Simpson et al., 2014). Such a poleward displacement
339 would be more remarkable during the seasonal transitions, when the surface temperatures
340 experience rapid changes and thus the jet location is more sensitive to the temperature
341 changes (Voigt and Shaw, 2016; Chen et al., 2020). Therefore, climate extremes during the
342 seasonal transitions should be emphasized, in addition to the extremes during the peak

343 warm or cold seasons that are currently under wide studies (Endo et al., 2021; Yang et al.,
344 2021; Zhou et al., 2022a). On the other hand, the present results suggest that climate
345 extremes in one region can be related to, or lead to, those in other regions, through the
346 atmospheric teleconnection patterns. It is widely believed that the climate extremes would
347 occur more frequently and their intensities would be enhanced under the global warming.
348 Therefore, the more frequent and stronger climate extremes would trigger the atmospheric
349 teleconnections more frequently, which in turn induce the climate extremes in remote
350 regions worldwide (Lau and Kim, 2012; Orsolini et al., 2015; Boers et al., 2019).

351

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357 ***Data Availability Statements.*** The gauge-based precipitation dataset over China can
358 be accessed at
359 [http://101.200.76.197/en/?r=data/detail&dataCode=SURF_CLI_CHN_MUL_DAY_CES](http://101.200.76.197/en/?r=data/detail&dataCode=SURF_CLI_CHN_MUL_DAY_CES_V3.0)
360 [_V3.0](http://101.200.76.197/en/?r=data/detail&dataCode=SURF_CLI_CHN_MUL_DAY_CES_V3.0). The precipitation dataset over India can be accessed at
361 https://www.imdpune.gov.in/Clim_Pred_LRF_New/Grided_Data_Download.html. ERA5
362 atmospheric reanalysis can be accessed at
363 [https://cds.climate.copernicus.eu/cdsapp#!/dataset/reanalysis-era5-pressure-](https://cds.climate.copernicus.eu/cdsapp#!/dataset/reanalysis-era5-pressure-levels?tab=overview)
364 [levels?tab=overview](https://cds.climate.copernicus.eu/cdsapp#!/dataset/reanalysis-era5-pressure-levels?tab=overview).

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